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## Role of iodine oxoacids in atmospheric aerosol nucleation

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## Role of iodine oxoacids in atmospheric aerosol nucleation

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## Abstract:

Iodic acid ( $\text{HIO}_3$ ) is known to form aerosol particles in coastal marine regions; but predicted nucleation and growth rates are lacking. Here, using the CERN CLOUD chamber, we find the nucleation rates of  $\text{HIO}_3$  particles are rapid, even exceeding sulfuric acid-ammonia rates under similar conditions. We find ion-induced nucleation involves  $\text{IO}_3^-$  and the sequential addition of  $\text{HIO}_3$ , and it proceeds at the kinetic limit below  $+10^\circ\text{C}$ . In contrast, neutral nucleation involves the repeated sequential addition of iodous acid ( $\text{HIO}_2$ ) then  $\text{HIO}_3$ , showing that  $\text{HIO}_2$  plays a key stabilizing role. Freshly formed particles are composed almost entirely of  $\text{HIO}_3$ , which drives rapid particle growth at the kinetic limit. Our measurements indicate that iodine oxoacid particle formation can compete with sulfuric acid in pristine regions of the atmosphere.

## One Sentence Summary

Iodic acid ( $\text{HIO}_3$ ) and iodous acid ( $\text{HIO}_2$ ) rapidly form new particles and can compete with sulfuric acid in pristine regions.

New particle formation plays an important role in radiative forcing of the climate. If particles survive to larger sizes, they influence climate directly by scattering light and indirectly by producing more than half of all cloud condensation nuclei (CCN) (1). However, new particle formation and aerosol-cloud interactions remain relatively poorly understood and constitute major uncertainties in determining Earth's equilibrium climate sensitivity with climate models (2). So far, only a few vapors that can form new particles under atmospheric conditions have been identified. They comprise sulfuric acid (3–7), methanesulfonic acid (3, 8), iodine species (9–11), highly oxygenated organic molecules (HOMs) (12) and, more recently, nitric acid (13). Under warm conditions or low vapor concentrations, acidic molecular clusters further require base vapors such as ammonia (6, 13, 14) or dimethylamine (15) to stabilize them against evaporation. Ions can play a comparable role in stabilizing nucleating acidic (6) or biogenic (12) particles.

Marine new particle formation is especially important as the ocean is vast, and marine clouds are highly sensitive to CCN since their number concentrations are low. Marine clouds are radiatively important since they have a high infrared emission and albedo in contrast with the dark ocean surface. Marine new particle formation has thus been a focus for aerosol-climate interactions and feedbacks for many years (16), mostly concerning dimethylsulfide and its oxidation products (3, 8, 16). Although nucleation of iodine oxides was first studied almost twenty years ago (9, 10), iodine particle formation is presently considered to have limited global significance (17) and remains relatively poorly understood.

The ocean surface is a major source of atmospheric iodine; hypoiodous acid (HOI) and molecular iodine ( $I_2$ ) are emitted by the reaction of gaseous ozone with aqueous iodide ( $I^-$ ) (18). Typical daytime emissions from the tropical Atlantic Ocean are  $7 \times 10^7$  HOI molecules  $cm^{-2} s^{-1}$  and  $7 \times 10^6$   $I_2$  molecules  $cm^{-2} s^{-1}$  (18), leading to daytime mixing ratios in the marine boundary layer of around 1 pptv HOI and 0.1 pptv iodine atoms (19). Land surfaces (vegetation and soils) emit comparable iodine fluxes as iodomethane ( $CH_3I$ ), producing boundary layer mixing ratios around 0.5 pptv and free tropospheric levels around 0.1 pptv (20). Iodine compounds have been found at widespread sites that include coastal regions (9, 10), Arctic sea ice (11, 21), the marine boundary layer (22), the lower (23) and upper free troposphere (24), and the stratosphere (25).

More recently, iodic acid ( $HIO_3$ ) has been shown to drive coastal marine new particle formation at Mace Head, Ireland (11), and intense iodine particle formation has been reported along China's coast (26). So far, however, controlled laboratory experiments under atmospheric conditions are lacking and so atmospheric observations of  $HIO_3$  cannot be connected with predicted particle formation and growth rates. A mechanism for  $HIO_3$  nucleation has been identified (11) but the ion-induced (charged) and neutral (uncharged) clusters were unseparated, and may involve different iodine compounds. Furthermore, iodine oxides ( $I_xO_y$ ) are presently considered to be the major species responsible for the growth and composition of iodic particles (9, 10, 27) although, once again, experimental measurements under atmospheric conditions are lacking.

## Nucleation and growth rates

Here we report iodine new particle formation experiments performed under marine boundary layer conditions in the CERN CLOUD chamber (Cosmics Leaving Outdoor Droplets; see Methods) between September 2017 and November 2019. The experiments were conducted at +10 °C and

-10 °C, at 34-73 % relative humidity (RH) and 20-46 ppbv ozone. We introduced molecular iodine from an evaporator (0.4-168 pptv, median  $I_2 = 6.4$  pptv) in the chamber.

The formation pathways of iodine oxoacids from iodine precursor vapors are not well understood, although computational studies assume that  $HO_x$  radicals are required to produce  $HIO_3$  (28, 29). We tested this hypothesis by using green light (528 nm) alone to photolyze  $I_2$  into iodine atoms. At full intensity, CLOUD's green light source photolyzes iodine vapor at a rate of  $7 \times 10^{-3} \text{ s}^{-1}$ , although most experiments were carried out at relative intensities of 10–20 %. Green light does not produce significant amount of  $HO_x$  since it cannot photolyze  $O_3$ . Nevertheless, we found that iodine atoms are rapidly oxidized in the presence of water vapor and ozone to produce both  $HIO_3$  and iodous acid ( $HIO_2$ ; Fig. S1). Iodine oxoacids can form from hydrated iodine atoms and iodine oxide radicals reacting with ozone, and from hydrolysis of  $I_xO_y$  (27). Under our experimental conditions, photolysis of  $I_2$  typically produces  $2 \times 10^5 \text{ I atom cm}^{-3} \text{ s}^{-1}$ , and roughly  $\sim 1$  pptv iodine monoxide (IO) radicals. The full range of conditions probed includes IO concentrations found in the open ocean marine boundary layer and remote free troposphere (18, 19, 23, 24; Table S1). Notably, at constant actinic flux,  $HIO_3$  acid increases linearly with iodine concentration, whereas  $HIO_2$  increases as the square root (Fig. S2). We speculate that iodine oxoacids form at CLOUD either from iodine radicals, e.g.  $I + H_2O + O_3 \rightarrow HIO_3 + OH$ , or the initial  $I_xO_y$  intermediates (e.g.  $I_2O_2 + H_2O \rightarrow HIO_2 + HOI$ ). Since ozone and water vapor are found throughout the troposphere, our findings imply that molecular iodine will produce iodine oxoacids even under cloudy daylight conditions with negligible ultraviolet irradiation.

We show in Fig. 1A our measured nucleation rates at 1.7 nm,  $J_{1.7}$ , versus the  $HIO_3$  concentration at +10 °C and -10 °C, and under three ionization conditions: neutral,  $J_n$  (ions eliminated from the chamber by a 20 kV  $\text{m}^{-1}$  electric field); galactic cosmic ray,  $J_{gcr}$  (boundary layer ion pair concentrations of around  $700 \text{ cm}^{-3}$ ) and beam enhanced,  $J_{beam}$  (ion pair concentrations around  $2500 \text{ cm}^{-3}$ , comparable to the upper free troposphere). The measurements were performed at contaminant ammonia levels near 3 pptv. The nucleation rates show a strong dependency on  $HIO_3$  concentration, charge and temperature. There is a large ion enhancement of the nucleation rate at +10 °C, whereas  $J_{gcr}$  and  $J_n$  are comparable at -10 °C. The nucleation rate increases rapidly as the temperature falls from +10 °C to -10 °C. For comparison, we include in Fig. 1A our previous measurements of  $J_{gcr}$  at 1.7 nm for sulfuric acid with 100 pptv  $NH_3$  (14), which show that the nucleation rate of iodine oxoacids exceeds that of  $H_2SO_4 - NH_3$  at the same acid concentrations.

In Fig. 1B we show the dependence on  $HIO_3$  concentration of the particle growth rates between 1.8 and 3.2 nm, at +10 °C and -10 °C. The growth rates of iodine oxoacid particles at +10 °C are identical to our measurements for  $H_2SO_4-NH_3$  particles between +5 °C and +20 °C at the same acid concentrations (30). The close agreement implies that the iodine oxoacid particles are growing at the dipole-dipole enhanced kinetic limit for  $HIO_3$ , with negligible evaporation at +10 °C or below. The measurements further indicate that  $HIO_3$  dominates the growth of iodine particles in this size range and above (since the Kelvin barrier falls with increasing size). This behavior is in marked contrast with previous studies which considered that iodine oxides are responsible for growth (10, 27). At -10 °C, the growth rate of  $HIO_3$  particles increases by a further factor 2. This faster growth exceeds the kinetic limit for the arrival rate of  $HIO_3$  monomers onto the particles and is attributed to additional growth from  $HIO_3$  molecular clusters – similar to the situation for  $H_2SO_4-$

DMA particles (31) – which provide a large pool of condensable material that is “hidden” from the HIO<sub>3</sub> monomer measurement.

Our nucleation and growth rate measurements (Fig. 1) indicate that HIO<sub>3</sub> concentrations above around  $3 \times 10^6 \text{ cm}^{-3}$  and  $1 \times 10^7 \text{ cm}^{-3}$  lead to rapid new particle formation at -10 °C and +10 °C, respectively. The survival probability of particles at low acid concentrations depends exponentially on the ratio of growth rate to condensation sink in the atmosphere, or wall-loss rate in a chamber. For CLOUD, the wall loss rate of sulfuric acid vapor is  $2.2 \times 10^{-3} \text{ s}^{-1}$  (30), which is comparable to the condensation sink in the pristine continental boundary layer. In clean marine regions of the boundary layer or in the upper free troposphere, the condensation sink is often as low as  $10^{-4}$ – $10^{-5} \text{ s}^{-1}$ . In such regions even lower HIO<sub>3</sub> concentrations will lead to sustained new particle formation and subsequent growth rates of a few 0.1 nm per hour. Under such conditions of extremely low HIO<sub>3</sub> concentrations and condensation sinks, it is likely that ions will be important to stabilize the embryonic clusters against evaporation, *i.e.* ion-induced nucleation will be the dominant mechanism.

### Particle formation mechanisms

In Fig. 2 we show mass defect plots of negatively charged clusters (Fig. 2A) and neutral clusters (Fig. 2B) containing up to five iodine atoms, measured during nucleation events. The event in panel A) is continued in Fig. S3 up to clusters containing twelve iodine atoms. Fig. S4 shows all identified peaks of the event in panel B), before summing over water molecules and charger ions as displayed in Fig. 2B. Further details of the charged and neutral clusters and their signal strengths are provided in Table S2.

For ion-induced nucleation (Fig. 2A), we observe a sequence of negatively-charged iodine clusters of the form  $(\text{HIO}_3)_{0-1}(\text{I}_2\text{O}_5)_n \cdot \text{IO}_3^-$ , involving sequential addition of HIO<sub>3</sub> followed by rapid dehydration of HIO<sub>3</sub>·HIO<sub>3</sub> pairs in the cluster to form I<sub>2</sub>O<sub>5</sub> (Fig. 3A), as previously observed at Mace Head (11). We find no nucleation occurs for positively charged iodine clusters (Fig. S5). This is clearly seen from the negative and positive charged particle spectra in Fig. S6; nucleation and growth only take place for negative particles. Almost all the negative particles have been neutralized by charge recombination before they reach 3 nm, and they continue to grow as mainly neutral particles. In Fig. 3A we provide a schematic representation of the mechanism for ion-induced iodic acid nucleation, interpreted from the mass defect plot (Fig. 2A).

Neutral nucleation of iodine oxoacid particles, however, proceeds by a different mechanism (Fig. 2B). We find that neutral HIO<sub>3</sub>/I<sub>2</sub>O<sub>5</sub> clusters are relatively weakly bound at these temperatures (none is detected above three iodine atoms). During neutral nucleation we observe repeated sequential addition of HIO<sub>2</sub> then HIO<sub>3</sub>. The nucleating cluster sequences are shown schematically in Fig. 3B. A certain fraction of HIO<sub>3</sub>·HIO<sub>2</sub> pairs in the cluster dehydrate to form iodine tetroxide (I<sub>2</sub>O<sub>4</sub>, Fig. 3C). Since the three sequences shown in Fig. 3C have comparable intensities, it indicates the formation rate of I<sub>2</sub>O<sub>4</sub> in the neutral clusters is comparable to the monomer collision rate (few  $10^{-3} \text{ s}^{-1}$ ). In contrast, the strict conversion of HIO<sub>3</sub>·HIO<sub>3</sub> pairs during ion-induced nucleation (Fig. 3A) shows that the formation rate of I<sub>2</sub>O<sub>5</sub> in the charged clusters is much faster than the monomer collision rate (few times  $10^{-2} \text{ s}^{-1}$ ).



Our measurements show that  $\text{HIO}_2$  plays a key role in stabilizing neutral  $\text{HIO}_3$  clusters. To assess this, we used quantum chemistry to compute the formation free energy of several molecular dimers involving  $\text{HIO}_3$  (Table S3 and Fig. S7). The most strongly bound dimer is  $\text{HIO}_3 \cdot \text{HIO}_2$  ( $-12.9 \text{ kcal mol}^{-1}$ ). On the other hand, the  $\text{HIO}_3 \cdot \text{HIO}_3$  ( $-7.7 \text{ kcal mol}^{-1}$ ) and  $\text{HIO}_3 \cdot \text{HOI}$  ( $-1.6 \text{ kcal mol}^{-1}$ ) dimers are both less stable. This argues in favor of the dominant  $\text{HIO}_3 \cdot \text{HIO}_2$  dimer shown in the 1 $\rightarrow$ 2 iodine step in Fig. 3B and not  $\text{HIO}_3 \cdot \text{HIO}_3$  at the current experimental conditions. Our measurements and quantum chemical calculations suggest that  $\text{HIO}_2$  stabilizes neutral  $\text{HIO}_3$  clusters with a 1:1 stoichiometry similar to that seen for ammonia stabilization of embryonic  $\text{H}_2\text{SO}_4$  clusters (6). For completeness, we note that, for neutral nucleation, we cannot exclude the direct addition of  $\text{I}_2\text{O}_4$  molecules from gas phase  $\text{I}_2\text{O}_4$ , which we measure at concentrations of around 1% of  $\text{HIO}_3$  (Table S2). We also note that  $\text{HIO}_3$  shows very weak affinity for pairing with a base ( $\text{HIO}_3 \cdot \text{NH}_3$  is  $-5.0 \text{ kcal mol}^{-1}$ ). Once the neutral particles exceed a critical size, they can continue to grow by condensation of  $\text{HIO}_3$  alone (Fig. 1B), and so growth is no longer limited by the lower concentrations of  $\text{HIO}_2$ .

### Ion-induced nucleation rate

To investigate ion-induced nucleation further, we have measured the collision rate coefficients,  $k_{i+1}$ , for each step in the process,  $\text{N}_i^- + \text{HIO}_3 \rightarrow \text{N}_{i+1}^-$ , where  $\text{N}_i^-$  represents a negatively charged cluster containing  $i$  iodine atoms (see He *et al.* (32), and Supplementary Materials for further details). The rate coefficients measured between neutral  $\text{HIO}_3$  monomers and charged clusters containing up to 11 iodine atoms are shown in Fig. 4A. Within measurement errors, we find the same  $\text{HIO}_3$  rate coefficient for each charged iodic cluster from the dimer to 11-mer, with a mean value  $[1.72 \pm 0.26 \text{ (stat.)} + 0.24/-0.21 \text{ (syst.)}] \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$  ( $1\sigma$  uncertainties). For comparison, we show the theoretical expectations for the rate coefficients for charged  $\text{HIO}_3$  clusters from average dipole orientation theory (ADO, red curve) (33) and its extensions: hard-sphere average dipole orientation theory (HSA, green curve) and surface charged capture theory (SCC, blue curve) (34). The latter SCC theory agrees closely with our measurements. We show in Fig. 4B the enhancement factors for charged versus neutral rate coefficients (ratios of the CLOUD measurements divided by the neutral rate coefficients, ignoring dipole enhancement). The enhancement during ion-induced nucleation, which averages 6.3, rapidly shepherds newly formed particles through the smallest size range where they are highly mobile and most vulnerable to scavenging loss, and contributes to the faster particle formation rate.

Our measurements of the individual collision rate coefficients for charged clusters containing up to 11 iodine atoms (Fig. 4A) – and their good agreement with theoretical expectations – show that ion-induced iodic acid nucleation proceeds at the kinetic limit and therefore is strictly a barrierless process rather than nucleation. Significant evaporation of any cluster in this range would lead to a tell-tale higher apparent rate coefficient for the previous cluster. We have confirmed kinetic formation of charged clusters in two further ways. Firstly, we have used a kinetic model to calculate the ion-induced component of  $J_{\text{gr}} (= J_{\text{in}} + J_n)$  at  $+10^\circ \text{C}$  (see Supplementary Materials for details) and find it is consistent with our experimental measurements (Fig. 1A). Secondly, we have calculated the reaction free energies and evaporation rates for several molecular clusters containing an  $\text{IO}_3^-$  ion (Table S4). Our calculations show that extremely low evaporation rates are expected for charged iodine clusters. They further indicate that the  $\text{HIO}_3 \cdot \text{HIO}_3 \cdot \text{IO}_3^-$  cluster is much less stable than  $\text{I}_2\text{O}_5 \cdot \text{IO}_3^-$ , supporting the sequence we observe experimentally in the first steps of

Fig. 3A. The  $\text{HIO}_3$  collision rate measurements in Fig. 4A confirm that ion-induced nucleation is indeed due to the sequential addition of  $\text{HIO}_3$  monomers and not, for example, to mixed accretion of  $\text{HIO}_3$  and  $\text{I}_2\text{O}_5$  molecules.

## Particle composition

The measurements presented in Fig. 1B provide strong evidence that  $\text{HIO}_3$  drives the growth of iodine particles above 1.8 nm. However, we have seen that additional iodine compounds play important roles during nucleation:  $\text{HIO}_2$  for neutral nucleation, and the formation of iodine oxides –  $\text{I}_2\text{O}_5$  and  $\text{I}_2\text{O}_4$  – in the charged and neutral clusters, respectively (Fig. 2). To what extent do these other iodine species contribute to particle growth at larger sizes?

To address this question, we have directly measured the composition of freshly nucleated iodine particles in the size range up to around 10 nm with a soft-ionization  $\text{Br}^-$ -FIGAERO (Filter Inlet for Gases and AEROSols) mass spectrometer. The majority (90%) of particles between 3 and 10 nm are neutral (Fig. S6). The FIGAERO collects particles on a Teflon filter for 30 min and then evaporates the sample with a controlled temperature ramp over the following 15 min at the inlet of a mass spectrometer. This measures the deposited mass of each chemical constituent of the particles and produces an individual “thermogram” of its volatility (evaporation temperature).

In Fig. 5A we show the evolution of particle size during a nucleation experiment. Particle sizes below 2.5 nm are measured by a PSM (Particle Size Magnifier) and above 4 nm by a nano-SMPS (Scanning Mobility Particle Sizer). In Fig. 5B we show the evolution of total particle volume derived from these measurements (blue circles). We also show in Fig. 5B the evolution of total  $\text{HIO}_3$  volume concentration in the particle phase, measured by the FIGAERO. The mass spectrum is dominated by the single channel,  $\text{HIO}_3$  (78% of the total mass, excluding water), with the rest being primarily  $\text{I}_2$ . We did not find a significant amount of  $\text{HIO}_2$  in the particle-phase, due to its low concentration. The mass spectrum shows that the freshly formed particles are composed almost entirely of  $\text{HIO}_3$ , and not  $\text{I}_2\text{O}_{4-5}$  as previously thought (9–11). This is independently confirmed by the close agreement seen in panel B) between the volume concentrations measured by the particle sizers and by the direct measurements of particulate  $\text{HIO}_3$  with the FIGAERO. We have confirmed that the evaporated  $\text{HIO}_3$  resulted from iodine acid in the particle phase – and not by thermal decomposition of other iodine compounds – by conducting laboratory calibrations. We nebulized iodine acid particles and then collected and analyzed them with the FIGAERO using the same procedures as for our experiments at CLOUD (see Supplementary Materials for details). The FIGAERO thermograms for the nebulized samples agree well with those obtained at CLOUD (Fig. S8).

## Climate implications

Sulfuric acid-ammonia nucleation is known to be important in relatively pristine environments such as the free troposphere (14) or the Antarctic coastal region (35). We show here that the nucleation rate of iodine oxoacids exceeds  $\text{H}_2\text{SO}_4$  –  $\text{NH}_3$  at the same acid concentrations. In pristine, cooler regions of the atmosphere,  $\text{HIO}_3$  concentrations above around  $10^6 \text{ cm}^{-3}$  will lead to copious new particle formation and sustained growth at a few times 0.1 nm per hour. The question then arises: are there pristine regions of the atmosphere where the concentrations of  $\text{HIO}_3$  are



comparable to or exceed sulfuric acid, or else ammonia is absent? For those regions,  $\text{HIO}_3$  could be the dominant source of new particles.

We show in Figs. S9 and S10 our global boundary layer measurements of  $\text{HIO}_3$  at ten sites. The conditions for abundant iodine new particle formation and rapid growth are frequently reached at mid-latitude coastal sites with marine algae, such as Mace Head and Helsinki, at coastal polar sites such as Villum and Ny Ålesund in the Arctic, or Neumayer in the Antarctic. Although measurements over the remote ocean are sparse, frequent new particle formation over the high Arctic pack ice has recently been reported, driven by  $\text{HIO}_3$  with little contribution from sulfuric acid (36).

The implications for the future are notable. Global iodine emissions have increased three-fold over the last 70 years and may continue to increase in the future as sea ice becomes thinner (37), and surface ozone increases (18). Any resultant increase of iodine CCN in the Arctic region could increase longwave radiative forcing from clouds and provide a positive feedback mechanism that accelerates the loss of sea ice. Iodine is also widespread in the free troposphere (24), where low temperatures, low condensation sinks, and high ion production rates from galactic cosmic rays favor iodine particle formation. Indeed, particulate iodate ( $\text{IO}_3^-$ ) has recently been observed near the tropopause at iodine mixing ratios around 0.1-0.5 pptv, and  $\text{IO}_3^-$  is the main iodine reservoir in the stratosphere (25) (further discussion is provided in the Supplementary Materials).

In summary, our study shows that iodine acid,  $\text{HIO}_3$ , is the major iodine species driving both nucleation and growth of iodine oxoacid particles in the boundary layer, and it remains as the dominant constituent in the particulate phase. We have further shown that iodic acid,  $\text{HIO}_2$ , plays a key role in neutral nucleation by stabilizing  $\text{HIO}_3$  clusters against evaporation – but is not important for particle growth at larger sizes. The efficacy of  $\text{HIO}_3$  to form new particles exceeds the  $\text{H}_2\text{SO}_4 - \text{NH}_3$  system at the same acid concentrations. Although atmospheric measurements remain limited, they are nevertheless sufficient to demonstrate the ubiquity of  $\text{HIO}_3$  and its potential to compete with sulfuric acid(-ammonia) particle formation in pristine regions of the atmosphere such as marine coasts, the Arctic boundary layer or the upper free troposphere.

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**Supplementary Materials:**

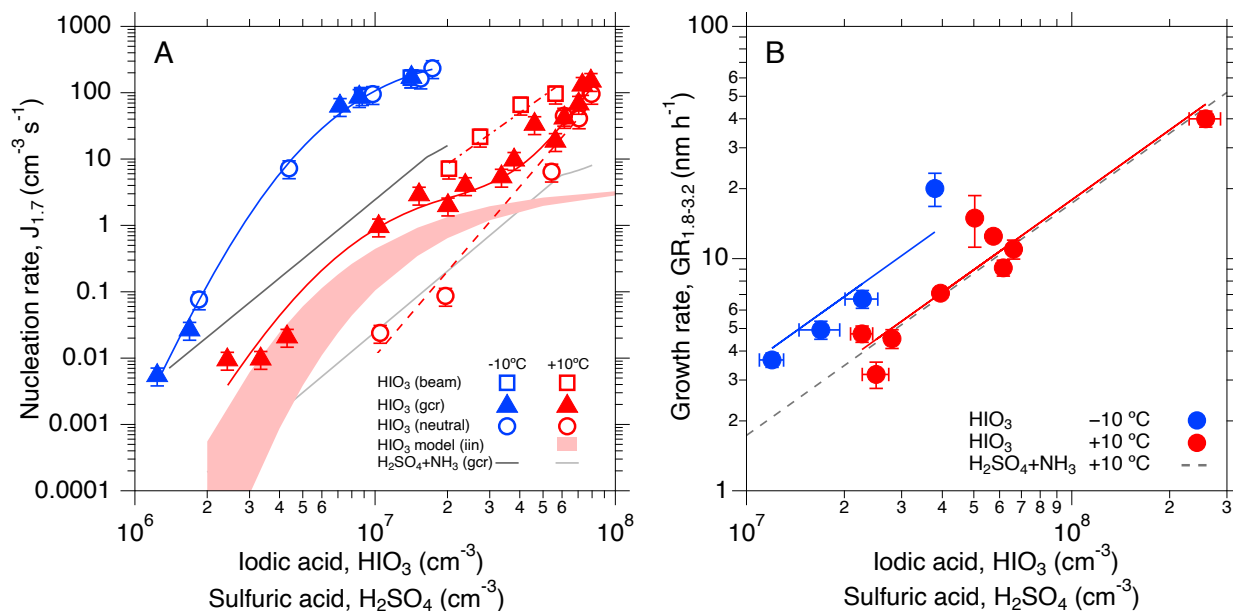
Materials and methods

Supplementary text

Tables S1-S4

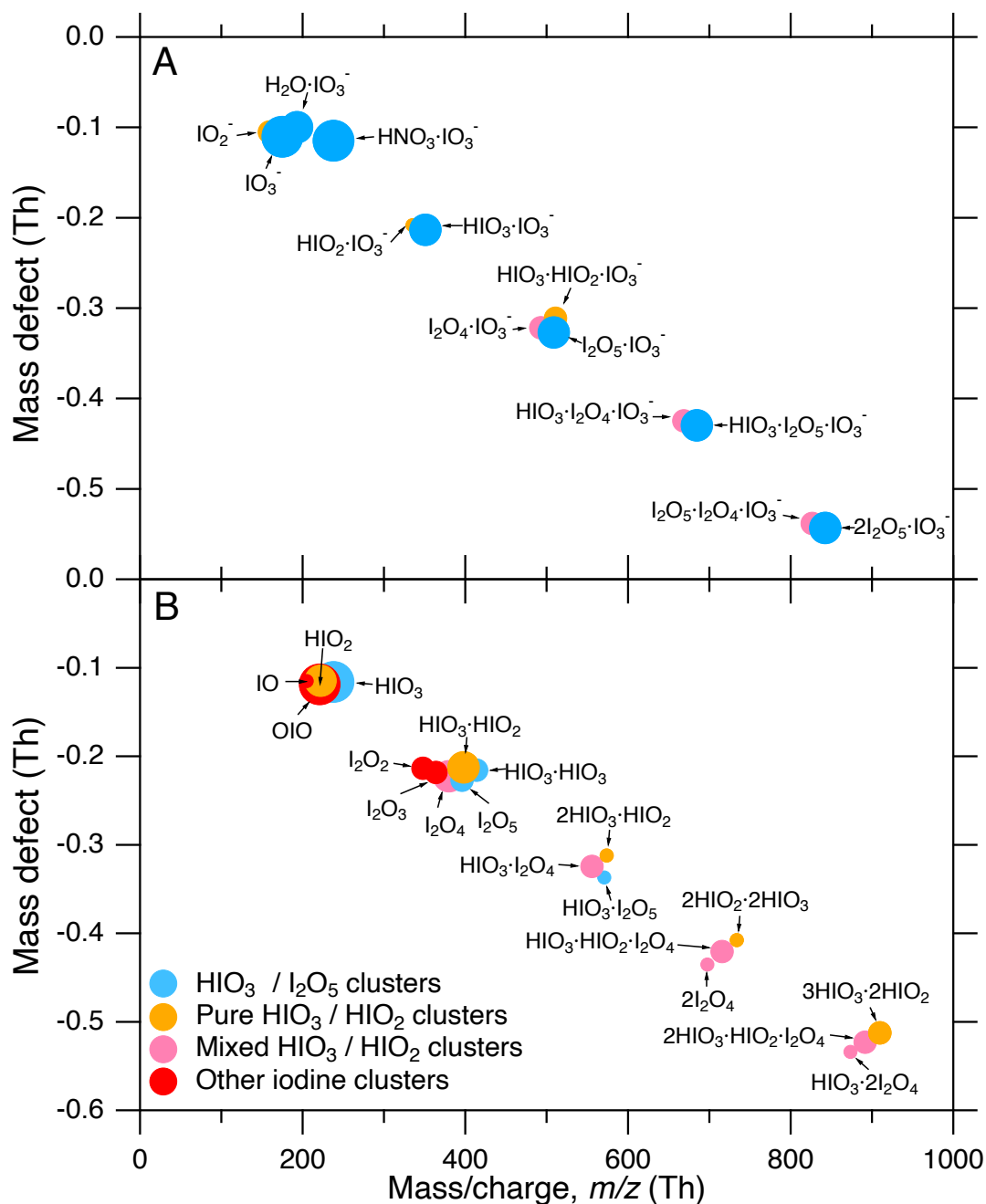
5      Figures S1-S10

References (39-81)



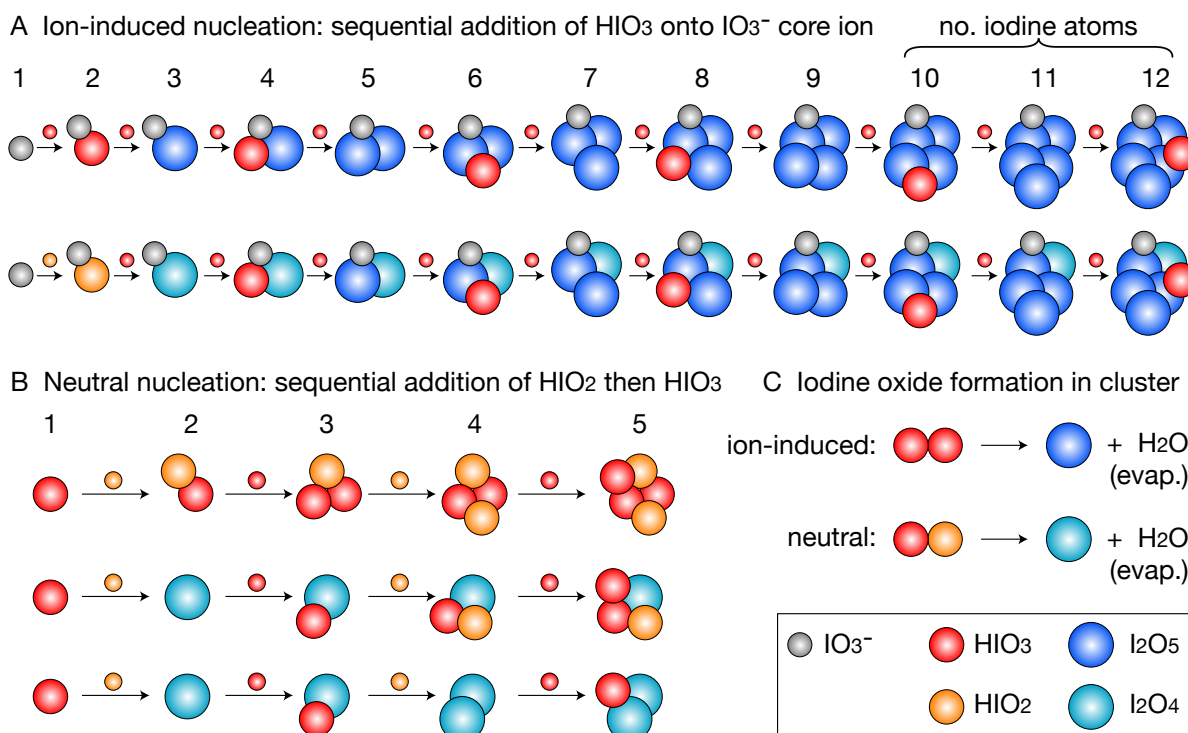
**Fig. 1. Nucleation and growth rates versus iodic acid concentration.** A) Nucleation rates at 1.7 nm diameter versus iodic acid concentration at +10 °C (red symbols and curves) and -10 °C (blue symbols and curves). Hollow circles show the nucleation rates for neutral conditions,  $J_n$ , solid triangles for gcr conditions,  $J_{gcr}$ , and hollow squares for beam conditions,  $J_{beam}$ . To guide the eye, the measurements are connected by approximate curves. The red band shows a kinetic model prediction for  $\text{HIO}_3$  ion-induced nucleation,  $J_{iin}$  ( $= J_{gcr} - J_n$ ), at +10 °C (see Supplementary Materials for further details). The lower and upper limits correspond, respectively, to zero and two  $\text{H}_2\text{O}$  molecules per iodine atom in the cluster. For comparison, the gcr nucleation rates measured for sulfuric acid with 100 pptv ammonia are shown at +10 °C (light grey curve) and -10 °C (dark grey curve) (14). B) Mean growth rates of particles (neutral, gcr and beam) between 1.8 nm and 3.2 nm diameter versus  $\text{HIO}_3$  concentration at +10 °C (filled red circles) and -10 °C (filled blue circles). For comparison, the dashed grey line shows the growth rates of  $\text{H}_2\text{SO}_4\text{-NH}_3$  particles measured at +10 °C (30). The bars in both panels represent  $\pm 1\sigma$  measurement errors. The experimental conditions are 36-44 ppbv  $\text{O}_3$ , 34-73% relative humidity (RH), 0.4-168 pptv  $\text{I}_2$  and an I atom production rate of  $4.4 \times 10^4$ - $1.5 \times 10^7 \text{ cm}^{-3} \text{s}^{-1}$ . An overall systematic scale error on the  $\text{HIO}_3$  concentration of -33%/+50% is not shown on the data points.



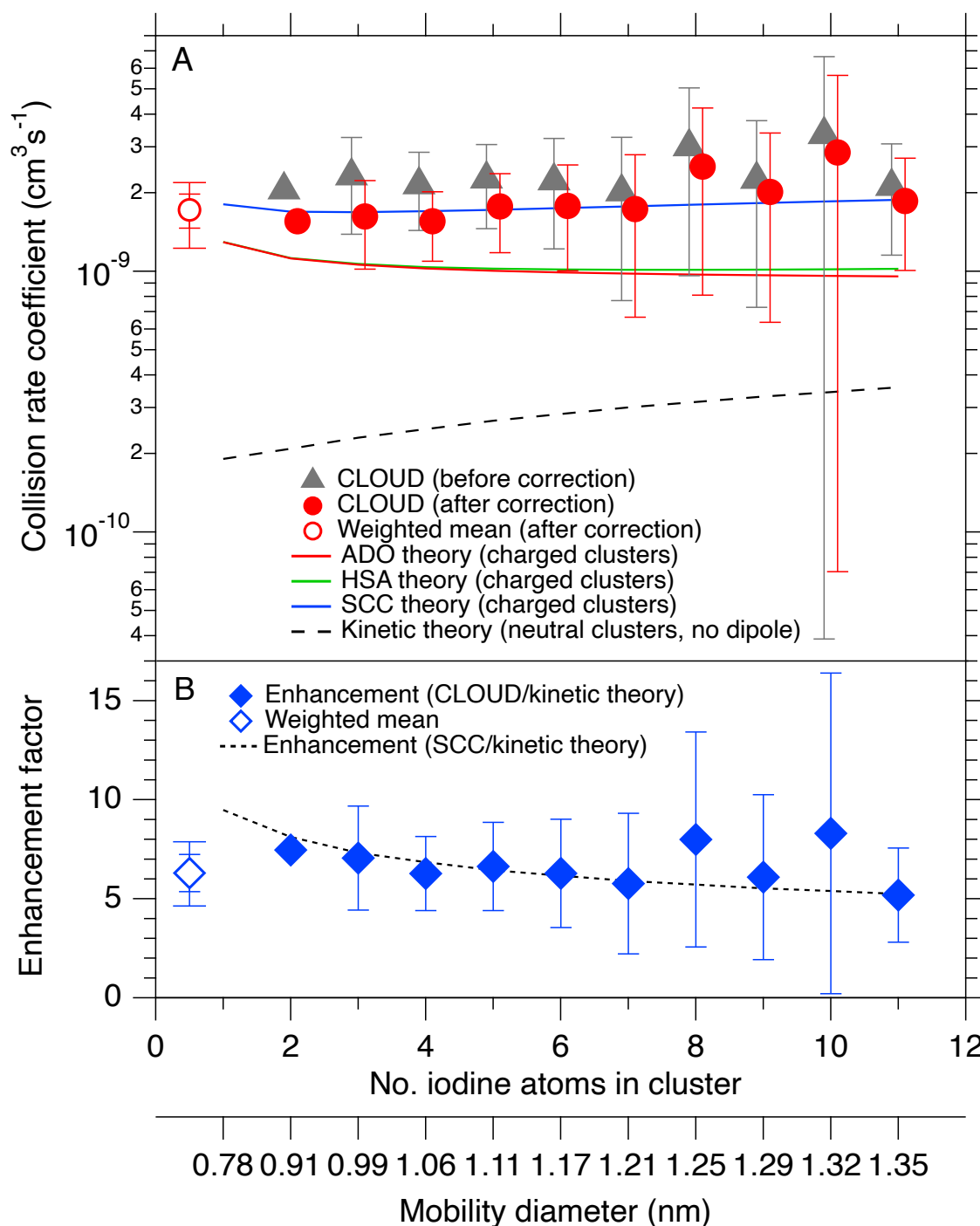


**Fig. 2. Charged and neutral mass defect plots during nucleation events.** Cluster mass defect (difference from integer mass) versus  $m/z$  of A) negatively charged and B) neutral clusters containing up to five iodine atoms during nucleation events. The experimental conditions are A) 36 ppbv  $\text{O}_3$ , 40% RH, +10 °C, 168 pptv  $\text{I}_2$  and  $1.5 \times 10^7$  I atom  $\text{cm}^{-3}\text{s}^{-1}$  and B) 46 ppbv  $\text{O}_3$ , 43% RH, +10 °C, 49 pptv  $\text{I}_2$  and  $2.4 \times 10^5$  I atom  $\text{cm}^{-3}\text{s}^{-1}$ . The event shown in panel A) is continued in Fig. S3 up to clusters containing twelve iodine atoms. In order to simplify panel B), water molecules and nitrate charger ions are ignored (Fig. S4 shows the same event where they are included). Charged clusters are measured with the APiTOF(-) and neutral clusters with the nitrate-CIMS

(preceded by an ion filter). We find no nucleation takes place for positively charged clusters (Figs. S5 and S6). Blue circles indicate clusters containing only  $\text{HIO}_3$  and  $\text{I}_2\text{O}_5$ . Orange circles indicate clusters containing only  $\text{HIO}_3$  and  $\text{HIO}_2$ . Pink circles indicate clusters containing  $\text{HIO}_3$ ,  $\text{HIO}_2$ ,  $\text{I}_2\text{O}_5$  and  $\text{I}_2\text{O}_4$ . Red circles indicate other iodine-containing neutral clusters. The area of the circle indicates signal strength on a logarithmic scale. Further details of the clusters and their signal strengths are provided in Table S2.

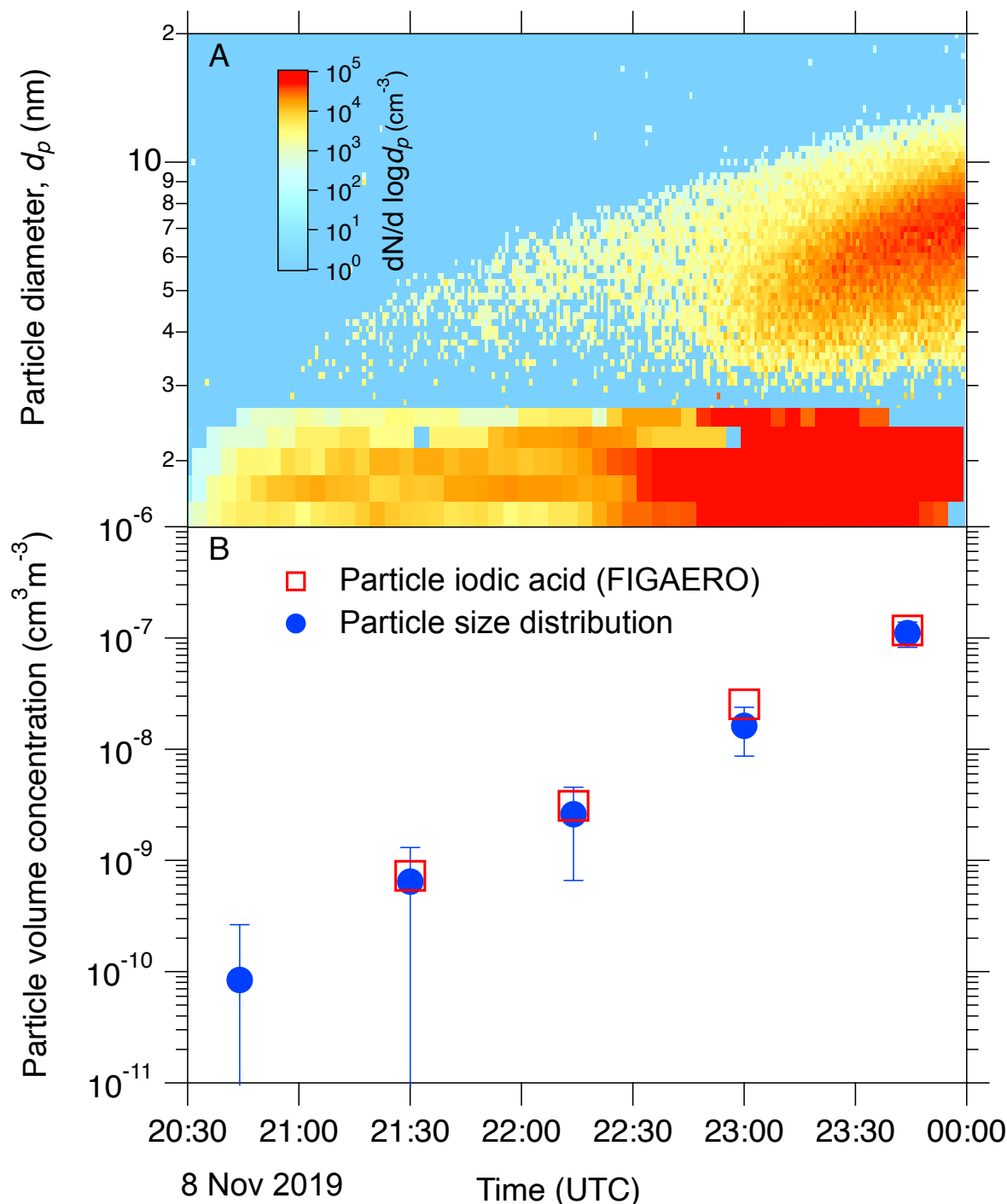


**Fig. 3. Nucleation mechanisms for iodine oxoacid clusters.** Schematic representations of the nucleation mechanisms for A) ion-induced (charged) and B) neutral (uncharged) iodine oxoacid clusters, interpreted from the mass defect plots. Ion-induced nucleation involves condensation of iodic acid (HIO<sub>3</sub>) alone onto an IO<sub>3</sub><sup>-</sup> ion, whereas neutral nucleation involves repeated stepwise condensation of iodous acid (HIO<sub>2</sub>) followed by iodic acid. Iodine oxide formation takes place in the clusters, as shown in panel C). Pairs of HIO<sub>3</sub> molecules always dehydrate to form I<sub>2</sub>O<sub>5</sub> in charged clusters (panel A). However, HIO<sub>3</sub> molecules do not form I<sub>2</sub>O<sub>5</sub> in neutral clusters but some may combine with HIO<sub>2</sub> and dehydrate to form I<sub>2</sub>O<sub>4</sub> (panel b). The relative intensities of the final neutral clusters in panel b) are (HIO<sub>3</sub>)<sub>3</sub>·(HIO<sub>2</sub>)<sub>2</sub> : (HIO<sub>3</sub>)<sub>2</sub>·HIO<sub>2</sub>·I<sub>2</sub>O<sub>4</sub> : HIO<sub>3</sub>·(I<sub>2</sub>O<sub>4</sub>)<sub>2</sub> = 0.38 : 0.46 : 0.16, indicating that the formation rate of I<sub>2</sub>O<sub>4</sub> in the neutral clusters is comparable to the monomer collision rate.



**Fig. 4. Collision rate coefficients for ion-induced iodic acid nucleation.** A) Collision rate (reaction rate) coefficients measured between neutral  $\text{HIO}_3$  monomers and charged clusters containing up to 11 iodine atoms. The experimental conditions are 20-41 ppbv  $\text{O}_3$ , 34-44% RH, +10 °C, 0.4-3.5 pptv  $\text{I}_2$  and  $(0.44\text{-}3.2) \times 10^5 \text{ I atom cm}^{-3} \text{s}^{-1}$ . The grey triangles are calculated from the 50% appearance times of a total of 8 experiments with  $(0.76\text{-}2.0) \times 10^7 \text{ cm}^{-3} \text{HIO}_3$ . The red

circles are the final experimental values after applying corrections from a kinetic model. The experimental points are horizontally shifted from integers to avoid overlaps. The solid curves show theoretical expectations for the charged collision rate coefficients from average dipole orientation theory (ADO, red curve (36)), hard-sphere average dipole orientation theory (HSA, green curve (37)) and surface charged capture theory (SCC, blue curve (37)). The expected collision rate coefficients between neutral monomers and neutral clusters, ignoring dipole-dipole interactions, are shown by the dashed black curve. Panel B) shows the measured enhancement factors for charged versus neutral collision rate coefficients (ratios of the corrected CLOUD measurements divided by the neutral collision rate coefficients). The black dotted line is the ratio of the SCC value to the neutral kinetic theory value. For both panels, the hollow markers show the weighted mean values from the trimer to 11-mer, with  $\pm 1\sigma$  errors indicating statistical without (inner caps) and with systematic errors (outer).



**Fig. 5. Evolution of particle size and chemical composition during iodic oxoacids nucleation.**

A) Evolution of the particle size measured by the PSM (below 2.5 nm) and nano-SMPS (above around 4 nm). The experimental conditions are 40 ppbv  $\text{O}_3$ , 40% RH, +10 °C, 8 pptv  $\text{I}_2$ ,  $(2.9-5.3) \times 10^5$  I atom  $\text{cm}^{-3} \text{s}^{-1}$  and  $(3.1-7.1) \times 10^7$   $\text{cm}^{-3}$   $\text{HIO}_3$ . The event is started by switching on green illumination (528 nm), and  $\text{HIO}_3$  is increased towards the end. B) Evolution of the particle volume concentration derived from i) the particle size distribution (blue circles) and ii) the  $\text{HIO}_3$  volume for particles collected and analyzed with the FIGAERO (hollow red squares). Particle



concentrations in the size range between 2.5 nm and 4 nm are obtained by interpolation between the PSM and nano-SMPS distributions and are verified by measurements of the total number concentrations above 2.5 nm threshold with the PSM. The FIGAERO collects particles on a Teflon filter for 30 min and then evaporates the sample with a controlled temperature ramp over the following 15 min at the inlet of the mass spectrometer. The FIGAERO data points are centered on the 30 min collection interval. The bars indicate  $\pm 1\sigma$  total errors. The FIGAERO mass spectrum shows that  $\text{HIO}_3$  dominates the particle composition (80% mass fraction). This is independently confirmed by the close agreement between the volume concentrations measured by the particle sizers and by the FIGAERO (panel A).

## Supplementary Materials for

### Role of iodine oxoacids in atmospheric aerosol nucleation

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Figs. S1 to S10  
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## 1 Materials and Methods

**The CLOUD experiment.** The experiments described in this study are conducted at the CLOUD atmospheric simulation chamber (Cosmics Leaving Outdoor Droplets) at CERN (European Centre for Nuclear Research), Geneva, Switzerland. The chamber has been described in detail elsewhere (e.g., Kirkby *et al.* (6) and Duplissy *et al.* (39)). The volume of the stainless-steel chamber is 26.1 m<sup>3</sup>. Ultra-pure synthetic air derived from cryogenic liquids (21% oxygen, 79% nitrogen) is continuously flowed into the chamber at 250-300 standard litre per minute (slpm). Magnetically-driven mixing fans at the top and bottom of the chamber establish near-homogeneous conditions within several minutes. The experiments are conducted under very clean conditions, with total light organics contamination below 150 pptv (6).

The synthetic air injected into the chamber is humidified with ultra-purified water. Ozone is generated in dry synthetic air under UV irradiation, and added to the main substitution flow. The injection of gaseous iodine into the chamber is provided from crystalline iodine (I<sub>2</sub>, Sigma-Aldrich, 99.999% purity) in a temperature-controlled evaporator to achieve levels of 0.4 to 168 pptv in the chamber. The injection lines are temperature stabilized and sulfinert-coated to minimize line conditioning effects. Trace gases are injected at the bottom of the chamber and dispersed in the chamber by the bottom fan.

The CLOUD chamber has unique control of ion concentrations. Electrodes installed in the chamber produce a 20 kV m<sup>-1</sup> electric field to sweep out ions in under 1 second (neutral conditions), so ions do not influence new particle formation or growth rates. When the electric clearing field is off, ions are produced by galactic cosmic rays (gcr conditions), allowing study of new particle formation under typical sea level ion concentrations. Additionally, irradiation of the chamber with the pion beam allows experiments at ion concentrations typical of the tropopause (beam conditions).

**Instrumentation.** A key factor limiting our understanding of iodine oxidation and new particle formation has so far been our inability to measure most of the gas-phase iodine species, as well as the difficulty of carrying out online-measurements of the composition of nucleation mode particles. In this study, we deployed a series of mass spectrometers to measure simultaneously the charged, neutral and particle phase iodine species. Among these instruments, a bromide chemical ionization atmospheric pressure interface time-of-flight mass spectrometer (bromide-CIMS) provides a high-sensitivity measurement of I<sub>2</sub>, and adapted Filter Inlet for Gases and AEROSols (FIGAERO)-CIMS (40) using bromide anions identifies the composition of nucleation mode particles.

**APi-TOF(+/-).** Atmospheric Pressure interface Time-Of-Flight mass spectrometers (APi-TOF, Aerodyne Inc.) are used to measure clusters charged by galactic cosmic rays or pions from the CERN beam (41). We deployed two APi-TOFs, one operating in positive mode and one in negative mode.

**Nitrate-CIMS.** The nitrate-CI-API-TOF (nitrate-CIMS) is an API-TOF coupled with a chemical ionization unit using nitric acid as the reagent gas. It is used extensively for detecting sulfuric acid, highly oxygenated organic molecules and iodic acid ( $\text{HIO}_3$ ). The instrument was originally described by Jokinen *et al.* (42), whereas details of the instrument used in the present study can be found in Kürten *et al.* (43). The nitrate-CIMS has an ion filter integrated in its sampling line in order to avoid confusion with ions and charged clusters from the CLOUD chamber; thus, it measures only neutral molecules and clusters in CLOUD.

**Bromide-CIMS.** The bromide-CI-API-TOF (bromide-CIMS) is an API-TOF coupled with a chemical ionization unit using dibromomethane ( $\text{CH}_2\text{Br}_2$ ) as the reagent gas. The  $\text{CH}_2\text{Br}_2$  is fed into the sheath flow of the inlet, under the illumination of a soft X-ray source, producing bromide anions ( $\text{Br}^-$ ). The  $\text{Br}^-$  ions are directed into the sample flow by a negative electric field, which then cluster with neutral molecules in the sample air. The bromide-CIMS in this study is mainly used to measure molecular iodine ( $\text{I}_2$ ), which is the precursor to produce  $\text{I}_x\text{O}_y$  and iodine oxoacids.

**CE-DOAS.** Gaseous molecular iodine concentrations are precisely measured with closed-path Cavity Enhanced Differential Optical Absorption Spectroscopy (CE-DOAS), using a green LED, high reflectivity mirrors (Advanced Thin Films), and a Czerny-Turner spectrometer with 0.5 nm optical resolution (Acton SP2150). The  $\text{I}_2$  limit of detection is 25 pptv for 1 min integration time, and 8 pptv for 10 min integration time. The optical path length in the instrument is around 20 km, determined by comparison of LED intensities in nitrogen and helium (44, 45). The accuracy for the CE-DOAS iodine time trace is estimated to be 20%, as a combination of cross section uncertainty and fit sensitivities.  $\text{I}_2$  concentrations are generally well below 10 pptv in most of the experiments but reach up to 1.7 ppbv during selected experiments for calibration purposes. Thereby, the CE-DOAS is used to calibrate the bromide-CIMS, which provides precise  $\text{I}_2$  concentrations at lower concentrations used in this study. A tight near-linear correlation to the bromide-CIMS time series is found and approximated by a quadratic fit. The calibration of the bromide-CIMS is estimated to be better than 30%.

**Light source.** We used a green light to photolyze  $\text{I}_2$  and to drive the iodine chemistry and subsequent new particle formation. The actinic fluxes for photolytic conversion of molecular iodine into iodine atoms are driven by an array of 48 green LEDs (light sabre 4, LS4), protruding into the chamber in the mid plane. The maximum total optical power output is 153 W, centered on 528 nm. This corresponds to a maximum iodine photolysis rate  $j = 7 \times 10^{-3} \text{ s}^{-1}$  and an  $\text{I}_2$  lifetime of 2.4 min. The photolysis rate is calculated from the decay rate of  $\text{I}_2$  at a fixed light intensity. Actinic fluxes are regulated by controlling the number of LEDs used and the set point of individual LED. Light fluxes are monitored by two photodiodes and a spectrometer. Finally, we calculate the iodine atom production rate by doubling the multiple of  $\text{I}_2$  concentration and iodine photolysis rate. Due to the high uncertainty in scaling the light fluxes to photolysis rates, we expect an overall a factor of three uncertainty in iodine atom production rate estimation.

**FIGAERO( $\text{Br}^-$ )-CIMS.** The FIGAERO( $\text{I}^-$ )-CIMS (40) is a widely used method to measure organic and inorganic aerosol composition. FIGAERO is a manifold inlet for API-TOF with two operating modes. In one mode, gases are directly sampled into a turbulent flow ion-molecule reactor while nucleated particles are concurrently collected on a PTFE (polytetrafluoroethylene) filter via a separate dedicated port. In the other mode, the filter is automatically moved into a pure

N<sub>2</sub> gas stream flowing into the ion molecule reactor while the N<sub>2</sub> is heated to evaporate the particles via temperature-programmed desorption. However, the use of iodide (I<sup>-</sup>) as the reagent ion would prevent it from being used to measure the composition of iodine-containing particles. Therefore, the FIGAERO used in this study uses CH<sub>2</sub>Br<sub>2</sub> as the reagent gas, since Br<sup>-</sup> has a high affinity for iodine-containing molecules. This is the first reported operation of a Br<sup>-</sup>-FIGAERO, which provides critical insight into the composition of iodine particles. The gas and particle phase sampling rates are 18 and 6 slpm, respectively. The particle collection period lasts 30 minutes and the detection limit of particulate HIO<sub>3</sub> is 3 ng m<sup>-3</sup>.

Particle-phase HIO<sub>3</sub> is quantified by a series of laboratory calibration experiments that collect suspended HIO<sub>3</sub> particles onto the FIGAERO filter over several time intervals (0, 15, 30 and 60 minutes) and analyze them with the same FIGAERO procedure as used at CLOUD. HIO<sub>3</sub> particles were constantly generated by nebulizing HIO<sub>3</sub> (≥ 99.5%, Sigma-Aldrich) water solution with an atomizer (TOPAS ATM 221), dried with a home-made diffusion dryer, and characterized with a TSI nano-SMPS. The generated particles have a geometric mean diameter of 14 nm and a total number concentration of 6.0×10<sup>3</sup> cm<sup>-3</sup>, comparable to those during the new particle formation experiments in CLOUD. To calculate the particle volume from HIO<sub>3</sub>, we assume the bulk density of HIO<sub>3</sub> (4.62 g cm<sup>-3</sup>). The log-scale slope between collected particle volume and particulate HIO<sub>3</sub> signal is better than 0.96. We note that, after going through the diffusion dryer, HIO<sub>3</sub> particles might not be completely effloresced, and so residual water content in the particles may cause a slight overestimate of the particle volume derived from FIGAERO particulate HIO<sub>3</sub> in Fig. 5B.

**PSM.** A scanning particle size magnifier (46) (PSM, Airmodus Ltd.) coupled to a condensation particle counter (CPC) is used to detect particles with a detection threshold between 1 and 3 nm. The PSM uses diethylene glycol as a working fluid to grow the particles before counting them by the CPC. The PSM is used to determine the nucleation rate at 1.7 nm ( $J_{1.7}$ ) and the number size distribution of particles between 1–3 nm.

**NAIS.** A Neutral cluster and Air Ion Spectrometer (NAIS) measures the mobilities and concentrations of the negative and positive clusters (47).

**DMA-train.** A differential mobility analyzer train (48) (DMA-train) measures the particle growth rates between 1.8 – 3.2 nm with high precision and sampling rate.

**Nano-SMPS.** A nano-scanning mobility particle sizer (nano-SMPS (49), TSI-3938) measures the particle size distribution between 4 and 65 nm. Particle concentrations in the size range between 2.5 and 4 nm are obtained by subtracting the total particle concentration above 4.0 nm measured by the nano-SMPS from the total particle concentration above 2.5 measured by the PSM. Particles between 2.5 nm and 4 nm are all considered 4 nm in the calculation of particle volume concentration in Fig. 5B, and it contributes less than 10% in the total volume concentration.

**Nucleation rates.** The nucleation rates of particles with diameter 1.7 nm and above are calculated using the same method as described in Dada *et al.* (50). Briefly, they are calculated according to the balance equation:

$$J = \frac{dN}{dt} + S_{dil} + S_{wall} + S_{coag} \text{ (cm}^{-3}\text{s}^{-1}\text{)}$$

where  $N$  is the particle number concentration above 1.7 nm;  $S_{dil}$ ,  $S_{wall}$  and  $S_{coag}$  are dilution, wall loss and coagulation losses, respectively. Typical values for  $S_{dil}$  and  $S_{wall}$  are  $1.6 \times 10^{-4} \text{ s}^{-1}$  and  $2.2 \times 10^{-3} \text{ s}^{-1}$ , respectively, while  $S_{coag}$  depends on the number and size distribution of particles in the chamber.

**Growth rates.** The particle growth rates are calculated by the 50% appearance time method as described in Lehtipalo *et al.* (51). Growth rates between 1.8–3.2 nm are calculated from the DMA-train.

**Quantum chemical calculations.** The initial conformer sampling of monomers was carried out in the Spartan 18 program using the MMFF force field (though most monomers in this study only had one conformer). Partial charges of different atoms in the monomers were then computed at the  $\omega$ B97X-D/aug-cc-pVTZ-PP level of theory by running a single-point calculation with the Pop=MKUFF keyword using the Gaussian 09/16 program (52). The monomer geometries and partial charges were then used by the ABCluster (53, 54) program to generate the initial structures for the dimer cluster. 200 initial structures were generated and the best 100 out of them were selected after the ABCluster procedure. During the ABCluster procedure, configurational sampling of the dimer clusters was performed using molecular mechanics (keeping the monomer structures rigid), with the intermolecular interactions defined by the CHARMM force field and the computed partial charges (55). Single point XTB calculations were further used to reduce the number of conformers for further analysis. Conformers within 7 kcal mol<sup>-1</sup> in relative electronic energy compared to the lowest-energy conformer were selected for further analysis. The initial monomer and cluster sampling, as well as the XTB energy calculations, used iodine atoms. However, the iodine atoms in the generated clusters were subsequently replaced by bromine, and the obtained clusters were optimized at the B3LYP/6-31+G\* level using Gaussian. The replacement of iodine by bromine was done because the (computationally very efficient) 6-31+G\* basis set has not been defined for iodine atoms. However, extensive test calculations (56) demonstrate that the geometries of I and Br-containing clusters are very similar. Optimized conformers within 6 kcal mol<sup>-1</sup> in relative electronic energy compared to the lowest-energy conformer were then selected for further calculations, and bromine atoms were again replaced back to iodine atoms. The structures were then re-optimized, and vibrational frequencies calculated, at the  $\omega$ B97x-D/aug-cc-pVTZ-PP (57, 58) level. Iodine pseudopotential definitions were taken from the EMSL basis set library (57). Similar procedures have been utilized in Hyttinen *et al.* (56). An additional coupled-cluster single-point energy correction was performed on the lowest-energy conformer calculated at the  $\omega$ B97x-D/aug-cc-pVTZ-PP level. The coupled cluster calculation was performed using the DLPNO-CCSD(T)/def2-QZVPP method with the ORCA program ver. 4.1.1 (59, 60). The polarizability and dipole moment of HIO<sub>3</sub> were also calculated at the  $\omega$ B97X-D/aug-cc-pVTZ-PP level.

**Kinetic model.** The calculated new particle formation rates at +10°C for the ion-induced case (HIO<sub>3</sub> model,  $J_{iin}$ ) in Fig. 1A are based on the numeric model presented by Kürten *et al.* (61) and Kürten (62). Despite the model was originally constructed to calculate neutral new particle formation rates, we adopted identical parameterizations presented in He *et al.* (32) and modified the model to simulate ion-induced iodic acid nucleation. It is shown in this study (figs. S3, S5 and



S6) and He *et al.* (32) that only negatively charged ion clusters grow and that this growth is mainly due the addition of neutral  $\text{HIO}_3$  molecules. Therefore, positively charged clusters are parameterized as sinks for negatively charged clusters. Additionally, since neutral clusters are shown to have a limited contribution to the ion-induced iodic acid nucleation (32), they are not considered in our simulations (contribution of neutral nucleation is not included in the simulations). In this way, we can clearly separate the contribution of ion-induced nucleation and neutral nucleation at the given experimental conditions. These assumptions lead to the following formulation of differential equations that are solved by the model. The balance equation for the positively charged monomer concentration ( $N_1^+$ ) is:

$$\frac{dN_1^+}{dt} = q - (k_{w,1} + k_{dil} + k_{rec} \cdot \sum_{i=1}^n N_i^-) \cdot N_1^+$$

where  $k_{w,1}$  is the wall loss rate for monomers,  $k_{dil}$  is the chamber dilution loss rate and  $k_{rec}$  is the recombination rate ( $2.2 \times 10^{-6} \text{ cm}^3 \text{ s}^{-1}$ ). The ion pair production rate,  $q$ , is taken as  $4.1 \text{ cm}^{-3} \text{ s}^{-1}$ . The negatively charged monomer concentration ( $N_1^-$ , i.e.,  $\text{IO}_3^-$ ) can be described by:

$$\frac{dN_1^-}{dt} = q - (k_{w,1} + k_{dil} + k_{rec} \cdot N_1^+ + k_{1,1} \cdot N_1^0) \cdot N_1^-$$

Here, it is important to note that the ion can grow due to the collision with a neutral acid molecule ( $N_1^0$ , i.e.,  $\text{HIO}_3$ ). The ion-neutral collision rate coefficient,  $k_{1,1}$ , for the present chemical system is  $1.7 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$ . The negatively charged cluster ion concentrations for  $i \geq 2$  ( $N_i^-$ , where  $i = 2$  corresponds to a negatively charged dimer) can be calculated according to:

$$\frac{dN_i^-}{dt} = k_{i-1,1} \cdot N_{i-1}^- \cdot N_1^0 - (k_{w,i} + k_{dil} + k_{rec} \cdot N_1^+ + k_{i,1} \cdot N_1^0) \cdot N_i^-$$

As the charged clusters grow at the kinetic limit, no evaporation rates are taken into account (32). Finally, the new particle formation rate for the ion-induced case,  $J_{iin}$ , is taken as the production term at the critical cluster size of 1.7 nm:

$$J_{iin} = k_{n_{crit}-1,1} \cdot N_{n_{crit}-1}^- \cdot N_1^0$$

In the model a fixed monomer acid concentration ( $N_1^0$ ) is used and the  $J_{iin}$  is taken as the steady-state value. The same ion-neutral collision rate coefficient ( $1.7 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1}$ , see above) is used for all cluster sizes as the rate coefficient is not expected to change significantly for charged clusters below 1.7 nm (e.g., Fig. 4 shows the rate coefficients of dimer to 11-mer). A value of  $4.62 \text{ g cm}^{-3}$  for the density (bulk density of iodic acid) and  $176 \text{ g mol}^{-1}$  for the molecular weight are used for  $\text{HIO}_3$  as the lower end of the simulation in Fig. 1A. Additionally, as Khanniche *et al.* (63) concluded that the dihydrates of  $\text{HIO}_3$  are stable below 310 K, we include a simulation with a value of  $2.86 \text{ g cm}^{-3}$  for the density and  $212 \text{ g mol}^{-1}$  for the molecular weight for  $\text{HIO}_3 + \text{two water molecules}$  ( $\text{HIO}_3 \cdot 2\text{H}_2\text{O}$ ) as the higher end of the simulation in Fig. 1A. The addition of water molecules increases the contributed volume of  $\text{HIO}_3$  monomer which in turn reduces the amount of  $\text{HIO}_3$  monomers needed to reach the critical cluster size of 1.7 nm. The simulation (red band in Fig. 1A) agrees well with our measured  $J_{gcr}$  within measurement uncertainties, below  $\text{HIO}_3$  of  $2 \times 10^7 \text{ cm}^{-3}$ , where  $J_n$  is negligible.

## 2 Collision rate coefficients

### 2.1 Calculation of collision rate coefficient from theoretical methods

There are three theoretical methods which are compared with our measurement data: the widely-used “average dipole orientation theory” (ADO theory) (33), the hard-sphere average dipole orientation theory (HSA theory), and the surface charge capture theory (SCC theory) (34). These methods are detailed in the original literature, and we only briefly compare them here. The ADO theory considers the thermal rotational energy of polar molecules, but it treats the charged cluster as a single point. The HSA theory extends the ADO theory by including the physical size of the charged cluster in the derivation, which in turn increases the predicted collision rate coefficients. This is particularly important for large charged clusters, while it has only a small effect for the charged clusters in this study ( $<1.5$  nm). Finally, the SCC theory assumes that the charge can freely move around the cluster. When collision occurs, the charge tends to move to the nearby surface of the cluster, which increases the effective capture radius and hence also the collision rate coefficient.

### 2.2 Calculation of apparent collision rate coefficient from measurement data

Detailed derivation of the appearance time method which we apply to the calculation of collision rate coefficients can be found in He *et al.* (32). Briefly, an APi-TOF operating in negative ion mode was deployed to measure the sequential growth of charged molecular clusters in ion-induced iodic acid nucleation experiments. The 50% appearance time method [APP50, (51)] was then applied to the time evolution of the charged clusters to obtain appearance times. The apparent collision rate coefficient between the  $(i+1)$ -mer cluster and  $\text{HIO}_3$  molecules is calculated according to

$$k_{i+1} = \frac{1}{[\text{HIO}_3]_{\text{avg}} \times (t_{i+1,50} - t_{i,50})}$$

where  $t_{i+1,50}$ ,  $t_{i,50}$  are the 50% appearance times of the  $(i+1)$ -mer and  $i$ -mer, respectively, and  $[\text{HIO}_3]_{\text{avg}}$  is the averaged  $\text{HIO}_3$  concentration during the time interval  $[t_{i+1,50}, t_{i,50}]$ . The results are shown in Fig. 4A as “CLOUD, before correction”. These values are also inferred as “apparent collision rate coefficients”.

### 2.3 Correction of the apparent collision rate coefficient

However, as detailed in He *et al.* (32), the deviation of the APP50 method neglected a number of processes that affect the ion distributions. These processes include ones that directly affect the full cluster population (coagulation, wall losses, dilution and evaporations), ones that only affect charged clusters (ion production in the chamber and ion-ion recombination), and the finite time resolution of the measurements of both physical and chemical properties. The Polar AND high-altitude Atmospheric research 520 (PANDA520) model was originally developed in He *et al.* (32) to simulate charged iodic acid cluster formation processes. In this study, we further develop a method to correct the apparent collision rate coefficient for the neglected processes.

In order to run the model, the collision rate coefficient based on surface charge capture (SCC) theory was input into the model. This is because the SCC theory produced the closest results to the apparent collision rate coefficient as shown in Fig. 4A. Then, the charged cluster distributions from the model output data were used to calculate the apparent collision rate coefficient based on the APP50 method. The apparent collision rate coefficient was then compared with the SCC theory values. Correction factors are thus derived for individual charged clusters as the ratio of the SCC predicted values to calculated apparent collision rate coefficients. The final measured collision rate coefficients are obtained by applying the correction factors to the apparent collision rate coefficients.

However, it needs to be noted that there are two sets of correction factors, since our experiments can either start from ion-free conditions or from ion-present conditions. First and the most common way to start an experiment was to switch on the green light to photolyze molecular iodine. This approach produces charged clusters and  $\text{HIO}_3$  from clean conditions. Some initial  $\text{IO}_3^-$  anions may be present owing to a low residual level of  $\text{HIO}_3$  that is almost always present in the chamber at the beginning of the experiments. Although the residual concentration is too low to grow charged clusters, it is sufficient to produce  $\text{IO}_3^-$  anions since  $\text{HIO}_3$  is a strong acid and thus an efficient proton donor. The residual  $\text{IO}_3^-$  can confound estimation of the appearance time for  $\text{IO}_3^-$ , and thereby, affect the estimation of the collision rate coefficient of  $\text{HIO}_3 \cdot \text{IO}_3^-$  with  $\text{HIO}_3$ . There were six experiments belong to this group.

The second way was to start an experiment from ion-free conditions. By turning the electric fields off, primary ions produced by gcr could survive and grow. The green light was always turned on to maintain a stable production of  $\text{HIO}_3$ . There were two experiments belong to this group. Only runs with the second method were used to calculate the collision rate coefficient of the dimer ( $k_2$ ). Due to this lack of statistics in the estimation of  $k_2$ , we do not determine a corresponding error bar.

We applied the correction factors produced by the PANDA520 model (on average 0.79) to the apparent collision rate coefficients derived from the measurements (denoted “CLOUD, before correction” in Fig. 4A) to calculate the final collision rate coefficient (“CLOUD, after correction”). Additionally, in order to rule out that the correction factors are sensitive to initial input values, we replaced the values calculated by the SCC theory with values from the ADO theory in the model. This yielded an average correction factor of 0.75—very close to the value obtained using the SCC rate coefficients as input. Thus, the values of the correction factors obtained indeed depend mainly on the loss processes, and not on the absolute values of the input collision rate coefficients (as long as reasonable values are used).

### 3 Measurement sites

#### 3.1 Hyytiälä

The SMEAR II (Station for Measuring Ecosystem-Atmosphere Relations) is a measurement station located in Southern Finland, 220 km northwest from Helsinki and about 60 km northeast from Tampere (61° 51' N, 24° 17' E). The measurement station is in a rural continental site surrounded by mostly Scots-pine forests (64).

### **3.2 Aboa**

The Finnish Antarctic research station (Aboa) is located on Basen Nunatak at Vestfjella mountains in Queen Maud Land, Eastern Antarctica ( $73^{\circ} 03' \text{ S}$ ,  $13^{\circ} 25' \text{ W}$ ). The measurement site is 480 m above sea level and 130 km south of the edge of sea ice (35).

### **3.3 Ny Ålesund**

The Gruebadet Observatory is located at about 50 m above sea level, 800 m southwest from the Ny-Ålesund village Svalbard ( $78^{\circ} 55' \text{ N}$ ,  $11^{\circ} 56' \text{ E}$ ).

### **3.4 Greenland**

Villum Research Station, Station Nord, is located in the northeastern corner of Greenland ( $81^{\circ} 36' \text{ N}$ ,  $16^{\circ} 40' \text{ W}$ ). It is on Prinsesse Ingeborg Halvø peninsula.

### **3.5 Beijing**

The Beijing site is located in the west campus of Beijing University of Chemical Technology ( $39^{\circ} 94' \text{ N}$ ,  $116^{\circ} 30' \text{ E}$ ) (65). The campus is surrounded by commercial properties and residential buildings. Thereby, the station represents a typical urban site.

### **3.6 Helsinki**

The SMEAR III station is an urban measurement site located in the Kumpula campus, University of Helsinki ( $60^{\circ} 12' \text{ N}$ ,  $24^{\circ} 58' \text{ E}$ ). It is about 4 km north-east from Helsinki city center (a coastal city), and on top of a hill (25 m above the sea level).

### **3.7 Mace Head**

The Mace Head Research Station is located in Connemara, County Galway, Western Ireland, on the coast of Atlantic Ocean ( $53^{\circ} 19' \text{ N}$ ,  $9^{\circ} 53' \text{ W}$ ) with regular tidal activity (10). The station is situated about 100 m from the shoreline.

### **3.8 Nanjing**

The Nanjing measurement station, Station for Observing Regional Processes of the Earth System, Nanjing University (SORPES-NJU), is located about 20 km northeast of Nanjing, Eastern China ( $32^{\circ} 07' \text{ N}$ ,  $118^{\circ} 57' \text{ E}$ ).

### **3.9 Réunion**

Réunion Island ( $21.2^{\circ} \text{ S}$ ,  $55.7^{\circ} \text{ E}$ ) is a volcanic island located in the southwestern part of the Indian Ocean. The measurement site was Maïdo-OPAR observatory (L'observatoire de physique de l'atmosphère de La Réunion), which is a modern research station located on top of an old volcanic caldera at 2160 m above the sea level. Maïdo-OPAR receives free tropospheric air masses during

the night-time and offers unique views on the southern hemispheric air masses. The station is part of many international research networks, e.g., GAW and ACTRIS, and hosts many permanent measurements on atmospheric composition.

### 3.10 Neumayer

The German Antarctica research station, Neumayer, is located on the Ekström ice shelf. The location of the site in 2016 was at 70° 40' N, 8° 16' E.

## 4 Implication of iodine oxoacid particle formation in the atmosphere

### 4.1 Pristine boundary layer

We are now able to assess with more confidence the global significance of iodine oxoacid particle formation. In coastal locations where active emissions occur (e.g., in Mace Head (10, 11)), iodine-containing species can contribute significantly to new particle formation and growth to cloud condensation nuclei size; both charged and neutral cluster formation processes can dominate the overall cluster formation processes depending on the  $\text{HIO}_3$  concentration. In polar regions, we observed significant geographical and temporal differences. At Aboa,  $\text{HIO}_3$  levels were modest and thus we expect that  $\text{HIO}_3$  has a minor role in new particle formation. At Villum station, neutral clusters formation dominated the iodine cluster formation processes because of the low temperature, while at Ny Ålesund and Neumayer III, both charged and neutral clusters potentially contributed to the cluster formation processes due to the relatively low temperature and moderate  $\text{HIO}_3$  concentration (fig S9). Although  $\text{HIO}_3$  was detected at Hyytiälä (boreal forest) and Réunion (at high altitude), the concentrations measured so far were too low for significant new particle formation.

### 4.2 Polluted boundary layer

Iodine particle formation is generally considered irrelevant for polluted urban environments. However, a feature of particular interest in our measurements is that we have measured noticeable  $\text{HIO}_3$  in all of the three city sites for the first time (Helsinki, Finland; Beijing and Nanjing, China). In Nanjing (an inland city, roughly 300 km from the coast), the  $\text{HIO}_3$  concentration was around  $10^5 \text{ cm}^{-3}$  for most of the time, but occasionally approaching  $10^6 \text{ cm}^{-3}$  (representing a growth rate of roughly  $0.2 \text{ nm h}^{-1}$ , at  $+10^\circ\text{C}$ ) in the measurement period. This suggests that  $\text{HIO}_3$  has a little contribution to the particle formation processes in Nanjing most of the time. In Beijing, the daily peak concentrations of  $\text{HIO}_3$  were above  $10^6 \text{ cm}^{-3}$  for almost the entire measurement period, and they were very often higher than  $2 \times 10^6 \text{ cm}^{-3}$ . This suggests that  $\text{HIO}_3$  could contribute at maximum  $0.4 \text{ nm h}^{-1}$  in the initial growth of newly formed particles in Beijing (figs. S9, S10). The seemingly small number cannot be ignored, since the typical particle growth rates in Beijing in was between  $1 - 3 \text{ nm h}^{-1}$  in August 2018 (66). Stolzenburg *et al.* concluded that the evaporation of sulfuric acid from particles above 2 nm is negligible and growth proceeds kinetically even with low ammonia concentration at temperatures between  $+5$  to  $+20^\circ\text{C}$  (30). Considering that  $\text{HIO}_3$  is much less volatile than  $\text{H}_2\text{SO}_4$ , the same conclusion could also apply. Thereby, our results hint that  $\text{HIO}_3$  could contribute to, despite not dominate, the initial growth of particles in Beijing in summer.

In Helsinki, we measured significant amount of  $\text{HIO}_3$  in August 2019. The daily peak concentrations were often higher than  $10^7 \text{ cm}^{-3}$ , and occasionally even reaching  $3 \times 10^7 \text{ cm}^{-3}$  at which ion-induced iodic acid nucleation starts to play a role. As shown in fig. S9, the approximate growth rates from  $\text{HIO}_3$  were often above  $1 \text{ nm h}^{-1}$ , and reached  $5 \text{ nm h}^{-1}$  at maximum. Considering that a mean growth rate in August at Helsinki was around  $6.5 \text{ nm h}^{-1}$  (67), the contribution from  $\text{HIO}_3$  in particle growth can be important at Helsinki in summer. Thereby,  $\text{HIO}_3$  can contribute to both the nucleation and initial growth during summertime in Helsinki.

While it is relatively well-known that iodine particle formation is important in pristine coastal environments, its contribution to growth in an inland city (Beijing) and the contribution to both nucleation and growth in a coastal city (Helsinki) were not expected. Sive *et al.* found terrestrial sources of methyl iodide from vegetation and soils (20), despite iodine species had been widely considered marine originated. These sources may explain why we have measured  $\text{HIO}_3$  ubiquitously at all the inland sites as shown in figs. S9 and S10.

### 4.3 Upper troposphere/low stratosphere

The widespread presence of iodine in the free troposphere has been reported repeatedly in the literature (23–25, 68, 69). Of particular note, iodine has been consistently observed in the vicinity of the tropopause (24, 68, 70, 71), where temperatures are lowest and the strong temperature dependence of iodine oxoacid nucleation found in this study is likely to make it more relevant. The upper troposphere/lower stratosphere (UTLS) is also where galactic cosmic ray fluxes are at their maximum (72), further enhancing efficient ion-induced iodic acid nucleation, which is found to proceed at the kinetic limit already at/below  $+10^\circ\text{C}$ . Additionally, large particles are sparse in the upper troposphere, decreasing the scavenging of condensable vapors such as  $\text{HIO}_3$ , as well as nano-clusters. All these conditions favor a potential contribution of iodine particle formation in these parts of atmosphere. New particle formation has been observed to be widespread in these regions over much of the world (73). The sharp conversion of gas-phase iodine to the particle phase, across the tropopause (25), consistent with the findings in this paper, suggests iodine likely participates in this widespread new particle formation. Indeed, particulate iodate ( $\text{IO}_3^-$ ) is observed near the tropopause, and  $\text{IO}_3^-$  is the main reservoir of iodine in the stratosphere (25). The rapid nucleation and growth from iodine species help to carry aerosols through size ranges that are especially susceptible to scavenging by other aerosols or grow them to sizes where later growth from other condensable matter becomes effective. This is especially important in the UTLS, where changes to the radiative budget can impact large scale atmospheric dynamics (74, 75). The observed threefold increase of surface iodine in recent decades (37, 76, 77) is expected to propagate to the UTLS (25) and therefore may exert a climate forcing via the nucleation mechanisms proposed in this study, among others.

The rather sparse observations of  $\text{HIO}_3$  around the world reveal the need for more dedicated field measurements in order to elucidate the role of iodine particle formation at the surface and aloft. Combined with global simulations, these measurements can aid to identify the role of iodine particle formation in past, present and future climate systems.

## 5 The identity of the measured $\text{HIO}_3$ signals

Gaseous  $\text{HIO}_3$  was first measured by Sipilä *et al.* in a coastal site in western Ireland (11). The measured gaseous  $\text{HIO}_3$  signal comprises a few different peaks in the mass spectrum measured by the nitrate-CIMS, i.e.,  $\text{IO}_3^-$ ,  $\text{HIO}_3 \cdot \text{NO}_3^-$ ,  $\text{HIO}_3 \cdot \text{HNO}_3 \text{NO}_3^-$  and a few hydrated forms of these charged clusters. The distribution of these anions depends mostly on the softness of the setting deployed by individual instruments. For example, a more fragmenting setting would increase the ratio of  $\text{IO}_3^-$  to  $\text{HIO}_3 \cdot \text{NO}_3^- / \text{HIO}_3 \cdot \text{HNO}_3 \text{NO}_3^-$ , and vice versa.

The authors have detailed several important experiments and discussed in the Supplementary Information on why the measured gaseous  $\text{HIO}_3$  was indeed present in the atmosphere rather than artificially produced by the nitrate-CIMS (11). For instance, the authors injected a substantial amount of ammonia through the ion source of the nitrate-CIMS which activated the surface of the ion source preventing nitric acid from entering the ion-molecule reaction chamber. As expected, the observed nitrate anions ( $\text{NO}_3^-$ ) almost dropped by an order of magnitude. However, the measured  $\text{HIO}_3$  clusters increased compared to the control experiment in which the ammonia was not injected. This was the direct evidence that the measured  $\text{HIO}_3$  clusters were not formed artificially from the nitrate anions by, for example,



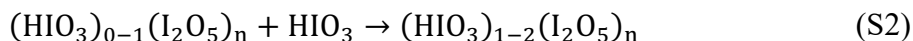
since otherwise the concentration of the measured  $\text{HIO}_3$  clusters would substantially decrease, rather than increase.

We further exclude this possibility in this study by deploying a nitrate-CIMS which was tuned to minimize the fragmentation in the ion optics of the instrument. We show a typical distribution of anions containing gaseous  $\text{HIO}_3$  in table S2. As can be seen, the  $\text{IO}_3^-$  signal only consists less than 6% of the total  $\text{HIO}_3$  signal, which in turn represents less than 6% of the reported gaseous  $\text{HIO}_3$  concentration. Even if the reaction (S1) could occur, the influence on our gaseous  $\text{HIO}_3$  measurements is still minor, not to say that this has already been proven not important in Sipilä *et al.* (11) at the first place.

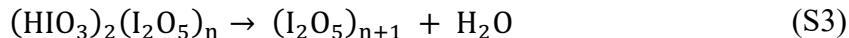
Another important question that has not been confirmed is the identity of the measured gaseous  $\text{HIO}_3$ .  $\text{HIO}_3$  has several isomers that can present in the atmosphere (28), and it has not been confirmed that the measured  $\text{HIO}_3$  is iodic acid ( $\text{HOIO}_2$ ). We confirm this by comparing the thermograms of the freshly formed small particles in the chamber and nebulized iodic acid – water solution in Fig. 5. As can be seen, the thermogram from the nebulized iodic acid–water solution shares identical features compared to the thermogram from freshly formed small particles in the chamber. This is a direct evidence that the measured  $\text{HIO}_3$  signal is gaseous iodic acid, since if the identity of the  $\text{HIO}_3$  signal is another isomer of iodic acid, the thermograms would not be identical.

## 6 Further remarks on the cluster formation mechanisms

Sipilä *et al.* (11) proposed that the general cluster formation mechanism for iodine-oxygen-hydrogen species in the atmosphere is



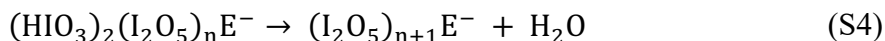




where  $n$  is an integer. The proposed mechanism consists of two major reaction types: first, the collision between  $\text{HIO}_3$  and iodine-containing clusters leads to cluster growth, and ultimately new particle formation. Additionally, the existence of two  $\text{HIO}_3$  in a cluster would lead to a chemical process in which the two  $\text{HIO}_3$  produce  $\text{I}_2\text{O}_5$  and  $\text{H}_2\text{O}$ , as is shown in reaction (S3).

However, in the past years, several issues have arisen which shed doubt on this mechanism. First of all, Khanniche *et al.* (78) calculated the rate coefficient for the reaction of two  $\text{HIO}_3$  forming one  $\text{I}_2\text{O}_5$  and an  $\text{H}_2\text{O}$ , and reported it to be  $5.56 \times 10^{-24} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$  at 300 K. This value, about 14 orders of magnitude below the gas-kinetic collision limit, would be far too slow to explain the fast cluster formation processes which have been observed in the atmosphere. Furthermore, their calculations indicate that the reverse reaction, in which an  $\text{I}_2\text{O}_5$  and a  $\text{H}_2\text{O}$  form two  $\text{HIO}_3$ , is much more favorable, and has a rate coefficient of  $1.03 \times 10^{-22} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$  at 300 K. Since the  $\text{H}_2\text{O}$  concentration is inevitably many orders of magnitude higher than the  $\text{HIO}_3$  concentration, gas-phase formation of  $\text{I}_2\text{O}_5$  from molecular  $\text{HIO}_3$  would not be expected to be important in our atmosphere.

Second, Passananti *et al.* (79) recently concluded that the instrumental setting is crucial in determining the extent of collision-induced cluster fragmentation, which in turn affects the mass spectra we observe. In addition, the charging process itself (even in the absence of high energy collisions) may be sufficient to catalyze or induce chemical reactions which would not occur for neutral molecules or clusters. For these reasons, Sipilä *et al.* (11) could not exclude the possibility that the reaction (S3) could happen in the instrument itself, rather than in the atmosphere. An alternative pathway to form the observed charged clusters would be:



where  $\text{E}^-$  is a core anion. It is thus unclear whether the chemical reaction of two  $\text{HIO}_3$  that leads to the formation of an  $\text{I}_2\text{O}_5$  and an  $\text{H}_2\text{O}$  is a natural process which happens in neutral clusters, or instead, it happens in the measurement instrument, either due to the charging, or the collision-induced cluster fragmentation, or some combination of both.

To minimize the collision-induced cluster fragmentation, the nitrate-CIMS instrument deployed in this study was a soft-tuned instrument. However, it should be noted that due to the limitation of the instrument itself, even a well-tuned instrument will still be affected by the fragmentation, though to a lesser extent. Notably, a series of water-containing nitrate clusters appear in the spectrum partly due to the soft-tuning. However, as the sum of the water-containing nitrate anions is less than 4% of the sum of nitrate anions in our instrument, water-containing nitrate anions only affect the species with significant abundance, e.g.,  $\text{HIO}_3$ . For measured neutral clusters which normally have a few counts or less than one count per second, the water-containing nitrate anions are unlikely to interfere with the detection.

One of the main results concerning the neutral cluster is that we have directly measured the  $2\text{HIO}_3 \cdot \text{NO}_3^-$  anion in the spectrum for the first time. This is the direct evidence that a neutral  $\text{HIO}_3$  dimer can be formed in the atmosphere without being completely converted to  $\text{I}_2\text{O}_5$ , which was

hypothesized by Sipilä *et al.* (11). The result also indicates that the conversion of  $\text{HIO}_3 \cdot \text{HIO}_3$  to  $\text{I}_2\text{O}_5$  does not happen instantaneously even in the presence of negative charge – at least in our soft-tuned instrument; a finite reaction rate may be expected. Additionally, the direct evaporation of  $\text{HIO}_3$  from iodine particles (Fig. 5) suggests that bulk  $\text{HIO}_3$  can indeed exist in iodine particles. However, we have not so far measured significant amount of  $\text{HIO}_2$  in particle phase, indicating either a negligible presence or that it might have been converted to, for example,  $\text{I}_2\text{O}_4$ .

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**Table S1. Comparison of CLOUD experiments with ambient conditions.** A summary of ambient measurements of iodine species can be found in Huang *et al.* (80). The pI is the iodine atom production rate, which is calculated either from I<sub>2</sub> photolysis or its reaction with NO<sub>3</sub> radicals. In the former case, the photolysis rate is assumed to be at mid-day sun condition, thus representing an upper limit. The I<sub>2</sub> photolysis rate at CLOUD is estimated from I<sub>2</sub> decay experiments at various light settings. We note that the iodine atom production rates from CLOUD and from ambient are not directly comparable as other halogen species and NO<sub>x</sub> can consume the produced iodine atoms and other initial iodine oxides in the ambient. This results in a higher yield of iodine oxoacids from iodine atoms in CLOUD than in the ambient. Nevertheless, the I<sub>2</sub>, HIO<sub>x</sub> and IO concentrations in the CLOUD experiments are comparable to those in the ambient.

Location	I <sub>2</sub> (pptv)	pI ( $\times 10^6$ atom cm <sup>-3</sup> s <sup>-1</sup> )	IO (pptv)
Mace Head day, low tide	20+	<sup>a</sup> 100+	4-10+
Mace Head day, high tide	5	<sup>a</sup> 30	2-7
Mace Head night (I <sub>2</sub> + NO <sub>3</sub> )	few 10	<sup>b</sup> 0.5	0.5-4
Open ocean	1	<sup>a</sup> 6	0.5-1
CLOUD median	6.4	<sup>c</sup> 0.2	<sup>d</sup> 1.2
CLOUD min-max	0.4 - 168	<sup>c</sup> 0.04-15	<sup>d</sup> 0.5-9.6

<sup>a</sup> Estimated from the I<sub>2</sub> concentration and an I<sub>2</sub> photolysis rate of 0.12 s<sup>-1</sup> with the mid-day sun (81).

<sup>b</sup> Estimated from the 10 pptv I<sub>2</sub>, 50 pptv NO<sub>3</sub> and a reaction rate coefficient of  $1.5 \times 10^{-12}$  cm<sup>3</sup> s<sup>-1</sup> with NO<sub>3</sub>.

<sup>c</sup> Estimated from I<sub>2</sub> concentrations and light fluxes of the reported experiments. The systematic uncertainty is a factor three.

<sup>d</sup> A systematic uncertainty of a factor of three is estimated for these measurements.



**Table S2. Molecular composition of charged and neutral clusters shown in Fig. 2.** The charged clusters (Fig. 2A) are produced by gcr ions in the CLOUD chamber. The neutral clusters (Figs. 2B and S4) are charged in the nitrate-CIMS by  $\text{NO}_3^-$  or  $\text{HNO}_3\cdot\text{NO}_3^-$  ions, as indicated in the table. Some of the clusters are ambiguous since we could not resolve whether  $\text{I}_2\text{O}_4$  or  $\text{I}_2\text{O}_5$  is present. In order to simplify Fig. 2B, all clusters with the same iodine molecular composition are summed into a single point, i.e., they are summed over all water molecules and nitrate charger ions.

<b>Charged clusters (Fig. 2A) (charged by gcr ions)</b>	<b>No. iodine atoms</b>	<b>Mass/charge, <math>m/z</math> (Th)</b>	<b>Signal (cps)</b>
<b>Pure <math>\text{HIO}_3</math> / <math>\text{HIO}_2</math> clusters (orange)</b>			
$\text{IO}_2^-$	1	158.895	0.331
$\text{HIO}_2\cdot\text{IO}_3^-$	2	334.791	0.021
$\text{HIO}_3\cdot\text{HIO}_2\cdot\text{IO}_3^-$	3	510.689	0.502
<b><math>\text{HIO}_3</math> / <math>\text{I}_2\text{O}_5</math> clusters (blue)</b>			
$\text{IO}_3^-$	1	174.890	40.423
$\text{H}_2\text{O}\cdot\text{IO}_3^-$	1	192.900	1.341
$\text{HIO}_3\cdot\text{NO}_3^-$	1	237.885	10.386
$\text{HIO}_3\cdot\text{IO}_3^-$	2	350.787	8.468
$\text{I}_2\text{O}_5\cdot\text{IO}_3^-$	3	508.673	5.259
$\text{HIO}_3\cdot\text{I}_2\text{O}_5\cdot\text{IO}_3^-$	4	684.570	2.946
$2\text{I}_2\text{O}_5\cdot\text{IO}_3^-$	5	842.457	2.062
<b>Mixed <math>\text{HIO}_3</math> / <math>\text{HIO}_2</math> clusters (pink)</b>			
$\text{I}_2\text{O}_4\cdot\text{IO}_3^-$	3	492.678	0.250
$\text{HIO}_3\cdot\text{I}_2\text{O}_5\cdot\text{IO}_3^-$	4	668.575	0.347
$\text{I}_2\text{O}_5\cdot\text{I}_2\text{O}_4\cdot\text{IO}_3^-$	5	826.461	0.430
<b>Neutral clusters (Figs. 2B and S4) (charged by nitrate anions)</b>	<b>No. iodine atoms</b>	<b>Mass/charge, <math>m/z</math> (Th)</b>	<b>Signal (cps)</b>
<b>Pure <math>\text{HIO}_3</math> / <math>\text{HIO}_2</math> clusters (orange)</b>			
$\text{HIO}_2\cdot\text{NO}_3^-$	1	221.885	1.638
$\text{HIO}_2\cdot\text{HNO}_3\text{NO}_3^-$	1	284.885	38.347
$\text{HIO}_3\cdot\text{HIO}_2\cdot\text{NO}_3^-$	2	397.788	79.746
$\text{HIO}_3\cdot\text{HIO}_2\cdot\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	415.798	0.708
$\text{HIO}_3\cdot\text{HIO}_2\cdot 2\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	433.810	0.795
$\text{HIO}_3\cdot\text{HIO}_2\cdot 3\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	451.820	0.632
$\text{HIO}_3\cdot\text{HIO}_2\cdot\text{HNO}_3\text{NO}_3^-$	2	460.784	10.120
$\text{HIO}_3\cdot\text{HIO}_2\cdot 4\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	469.831	0.564
$\text{HIO}_3\cdot\text{HIO}_2\cdot\text{H}_2\text{O}\cdot\text{HNO}_3\text{NO}_3^-$	2	478.794	0.299
$\text{HIO}_3\cdot\text{HIO}_2\cdot 5\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	487.843	0.408
$\text{HIO}_3\cdot\text{HIO}_2\cdot 2\text{H}_2\text{O}\cdot\text{HNO}_3\text{NO}_3^-$	2	496.808	0.380
$\text{HIO}_3\cdot\text{HIO}_2\cdot 6\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	505.853	0.289
$\text{HIO}_3\cdot\text{HIO}_2\cdot 3\text{H}_2\text{O}\cdot\text{HNO}_3\text{NO}_3^-$	2	514.816	0.346
$\text{HIO}_3\cdot\text{HIO}_2\cdot 7\text{H}_2\text{O}\cdot\text{NO}_3^-$	2	523.865	0.254
$\text{HIO}_3\cdot\text{HIO}_2\cdot\text{IO}_3^-$	3	510.691	0.344
$2\text{HIO}_3\cdot\text{HIO}_2\cdot\text{NO}_3^-$	3	573.688	0.455
$2\text{HIO}_3\cdot 2\text{HIO}_2\cdot\text{NO}_3^-$	4	733.592	0.517
$2\text{HIO}_3\cdot 2\text{HIO}_2\cdot\text{IO}_3^-$	5	846.491	1.560
$3\text{HIO}_3\cdot 2\text{HIO}_2\cdot\text{NO}_3^-$	5	909.488	0.306

<b>HIO<sub>3</sub> / I<sub>2</sub>O<sub>5</sub> clusters (blue)</b>			
IO <sub>3</sub> <sup>-</sup>	1	174.889	194.174
HIO <sub>3</sub> ·NO <sub>3</sub> <sup>-</sup>	1	237.884	2472.274
HIO <sub>3</sub> ·H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	255.895	23.859
HIO <sub>3</sub> ·2H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	273.906	11.635
HIO <sub>3</sub> ·3H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	291.916	12.686
HIO <sub>3</sub> ·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	300.880	782.010
HIO <sub>3</sub> ·4H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	309.927	11.424
HIO <sub>3</sub> ·H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	318.891	3.730
HIO <sub>3</sub> ·5H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	327.938	7.137
HIO <sub>3</sub> ·2H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	336.902	4.041
HIO <sub>3</sub> ·6H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	345.948	6.362
HIO <sub>3</sub> ·3H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	354.913	4.141
HIO <sub>3</sub> ·7H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	363.959	5.349
HIO <sub>3</sub> ·4H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	372.923	3.276
HIO <sub>3</sub> ·8H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	381.970	4.057
HIO <sub>3</sub> ·5H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	390.934	3.575
HIO <sub>3</sub> ·9H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	399.981	3.741
HIO <sub>3</sub> ·6H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	408.945	3.147
HIO <sub>3</sub> ·10H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	417.992	2.500
HIO <sub>3</sub> ·7H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	426.956	2.027
HIO <sub>3</sub> ·11H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	436.002	2.437
HIO <sub>3</sub> ·8H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	444.966	1.645
HIO <sub>3</sub> ·12H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	454.013	2.259
HIO <sub>3</sub> ·9H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	462.978	1.349
HIO <sub>3</sub> ·13H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	472.024	1.306
HIO <sub>3</sub> ·10H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	480.989	1.101
HIO <sub>3</sub> ·14H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	490.034	1.172
HIO <sub>3</sub> ·11H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	498.998	0.806
HIO <sub>3</sub> ·15H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	508.046	0.826
HIO <sub>3</sub> ·12H <sub>2</sub> O·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	1	517.011	0.599
HIO <sub>3</sub> ·16H <sub>2</sub> O·NO <sub>3</sub> <sup>-</sup>	1	526.057	0.677
I <sub>2</sub> O <sub>5</sub> ·NO <sub>3</sub> <sup>-</sup>	2	395.773	6.398
2HIO <sub>3</sub> ·NO <sub>3</sub> <sup>-</sup>	2	413.784	3.811
I <sub>2</sub> O <sub>5</sub> ·IO <sub>3</sub> <sup>-</sup>	3	508.675	0.423
<b>Mixed HIO<sub>3</sub> / HIO<sub>2</sub> clusters (pink)</b>			
I <sub>2</sub> O <sub>4</sub> ·NO <sub>3</sub> <sup>-</sup>	2	379.777	3.654
I <sub>2</sub> O <sub>4</sub> ·HNO <sub>3</sub> NO <sub>3</sub> <sup>-</sup>	2	442.773	34.927
HIO <sub>3</sub> ·I <sub>2</sub> O <sub>4</sub> ·NO <sub>3</sub> <sup>-</sup>	3	555.676	1.887
2I <sub>2</sub> O <sub>4</sub> ·NO <sub>3</sub> <sup>-</sup>	4	697.565	0.411
HIO <sub>3</sub> ·HIO <sub>2</sub> ·I <sub>2</sub> O <sub>4</sub> ·NO <sub>3</sub> <sup>-</sup>	4	715.579	1.583
2I <sub>2</sub> O <sub>4</sub> ·IO <sub>3</sub> <sup>-</sup>	5	810.470	0.435
HIO <sub>3</sub> ·HIO <sub>2</sub> ·I <sub>2</sub> O <sub>4</sub> ·IO <sub>3</sub> <sup>-</sup>	5	828.480	1.985
HIO <sub>3</sub> ·2I <sub>2</sub> O <sub>4</sub> ·NO <sub>3</sub> <sup>-</sup>	5	873.466	0.339
2HIO <sub>3</sub> ·HIO <sub>2</sub> ·I <sub>2</sub> O <sub>4</sub> ·NO <sub>3</sub> <sup>-</sup>	5	891.477	0.297
<b>Other iodine clusters (red)</b>			

$\text{IO} \cdot \text{NO}_3^-$	1	204.885	0.707
$\text{OIO} \cdot \text{NO}_3^-$	1	220.881	145.760
$\text{OIO} \cdot \text{HNO}_3\text{NO}_3^-$	1	283.877	84.952
$\text{I}_2\text{O}_2 \cdot \text{NO}_3^-$	2	347.786	4.121
$\text{I}_2\text{O}_3 \cdot \text{NO}_3^-$	2	363.782	6.426

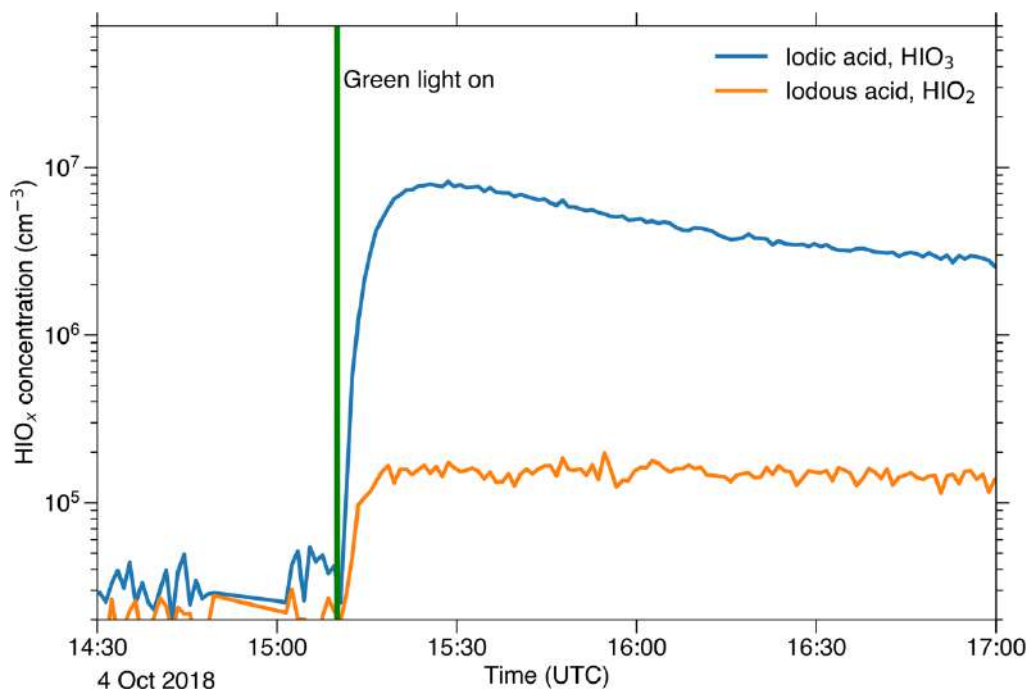
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**Table S3. Cluster formation free energies for neutral clusters.** The quantum chemical calculations are based on (DLPNO-CCSD(T)/def2-QZVPP// $\omega$ B97xD/aug-cc-pVTZ-PP level) at 283.15 K and 1 atm. The optimized geometries at the  $\omega$ B97xD/aug-cc-pVTZ-PP level are shown in fig. S7. The most stable dimer with  $\text{HIO}_3$  is  $\text{HIO}_3 \cdot \text{HIO}_2$ .

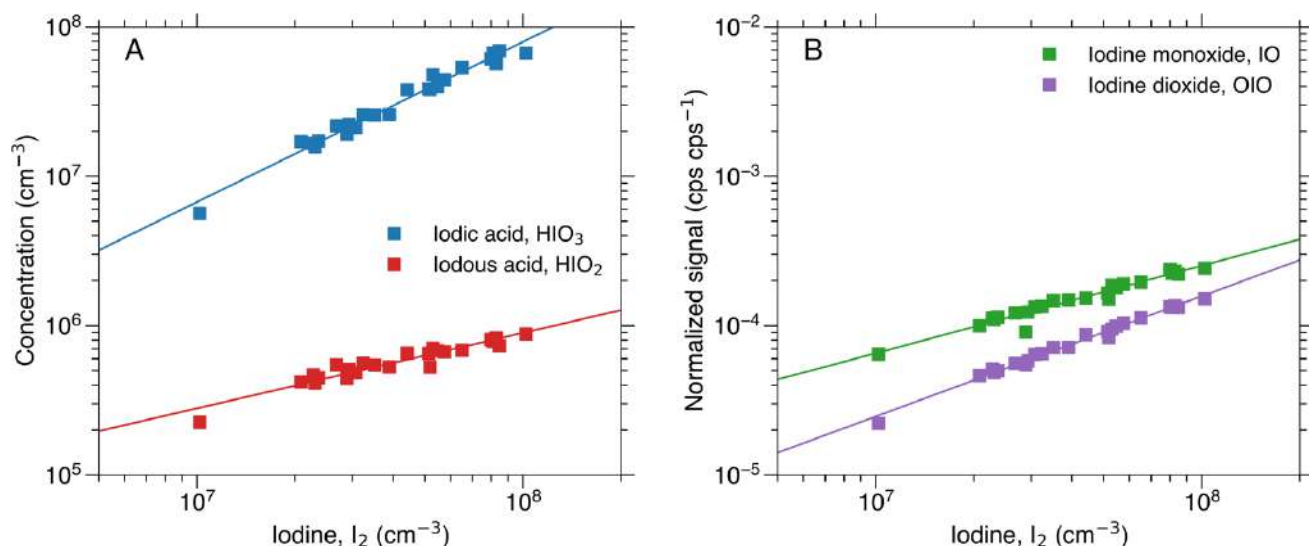
Cluster	Formation free energy (kcal mol <sup>-1</sup> )
$\text{HIO}_3 \cdot \text{HIO}_3$	-7.7
$\text{HIO}_3 \cdot \text{HIO}_2$	-12.9
$\text{HIO}_2 \cdot \text{HIO}_2$	-13.1
$\text{HIO}_3 \cdot \text{I}_2\text{O}_5$	-9.3
$\text{HIO}_3 \cdot \text{NH}_3$	-5.0
$\text{HIO}_3 \cdot \text{HOI}$	-1.6

**Table S4. Reaction free energies and calculated evaporation rates for charged clusters.** The quantum chemical calculations are based on (DLPNO-CCSD(T)/def2-QZVPP// $\omega$ B97xD/aug-cc-pVTZ-PP level) at 283.15 K and 1 atm.

Evaporation process	Reaction Free Energy (kcal mol <sup>-1</sup> )	Evaporation rate (s <sup>-1</sup> )
$\text{I}_2\text{O}_5 \cdot \text{IO}_3^- \rightarrow \text{I}_2\text{O}_5 + \text{IO}_3^-$	31.1	$4.2 \times 10^{-14}$
$\text{HIO}_3 \cdot \text{I}_2\text{O}_5 \cdot \text{IO}_3^- \rightarrow \text{I}_2\text{O}_5 \cdot \text{HIO}_3 + \text{IO}_3^-$	43.5	$1.1 \times 10^{-23}$
$\text{HIO}_3 \cdot \text{I}_2\text{O}_5 \cdot \text{IO}_3^- \rightarrow \text{I}_2\text{O}_5 \cdot \text{IO}_3^- + \text{HIO}_3$	21.6	$9.6 \times 10^{-07}$
$\text{HIO}_3 \cdot \text{I}_2\text{O}_5 \cdot \text{IO}_3^- \rightarrow \text{HIO}_3 \cdot \text{IO}_3^- + \text{I}_2\text{O}_5$	27.1	$4.6 \times 10^{-11}$
$\text{HIO}_3 \cdot \text{IO}_3^- \rightarrow \text{HIO}_3 + \text{IO}_3^-$	25.6	$7.2 \times 10^{-10}$
$\text{HIO}_3 \cdot \text{HIO}_3 \cdot \text{IO}_3^- \rightarrow \text{HIO}_3 \cdot \text{IO}_3^- + \text{HIO}_3$	17.4	$1.6 \times 10^{-03}$
$\text{HIO}_3 \cdot \text{HIO}_3 \cdot \text{IO}_3^- \rightarrow \text{HIO}_3 \cdot \text{HIO}_3 + \text{IO}_3^-$	33.6	$3.4 \times 10^{-16}$

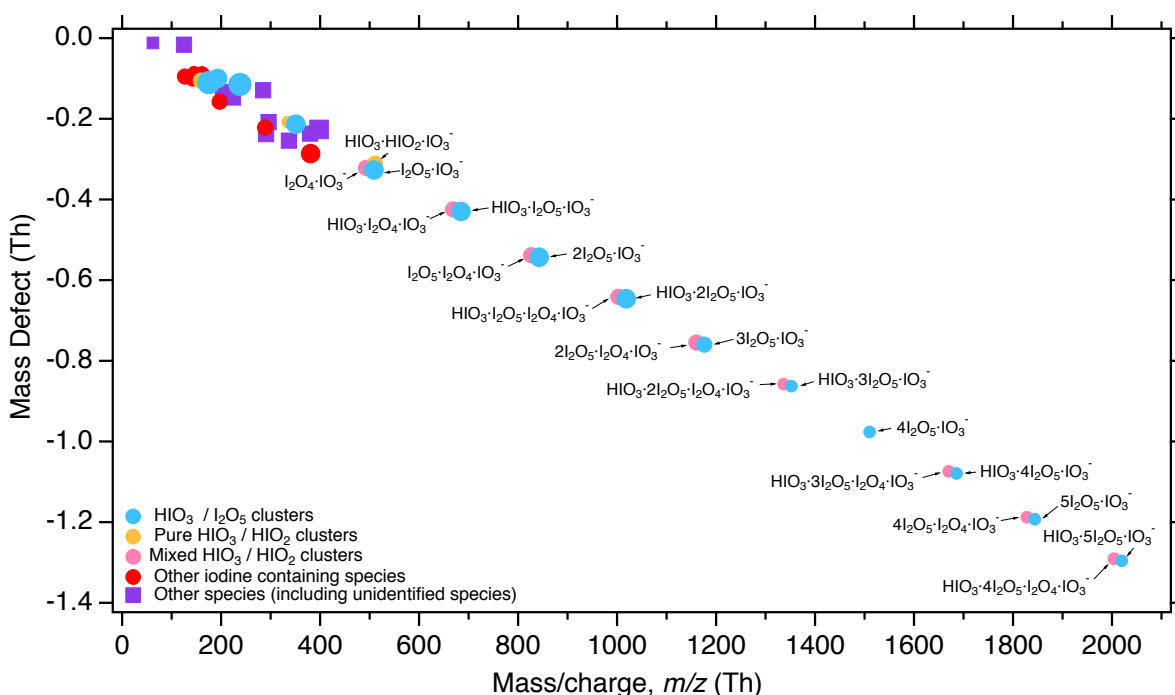


**Fig. S1. Formation of iodine oxoacids without  $\text{HO}_x$  radicals.** Gas-phase  $\text{HIO}_3$  and iodous acid ( $\text{HIO}_2$ ) measured in the CLOUD chamber before and after switching on green light (528 nm). The experimental conditions are 7 pptv  $\text{I}_2$ , I atoms production rate of  $6.2 \times 10^4 \text{ cm}^{-3} \text{ s}^{-1}$ , 40 ppbv ozone, 69% RH and  $-10^\circ\text{C}$ . Green light photolyzes  $\text{I}_2$  to iodine atoms but does not photolyze ozone and so  $\text{HO}_x$  is absent. This demonstrates that  $\text{HIO}_x$  can be produced by oxidation with ozone in the absence of  $\text{HO}_x$ . The measured  $\text{HIO}_2$  concentration assumes the same mass-spectrometer calibration factor as for  $\text{HIO}_3$ , and so represents a lower limit.

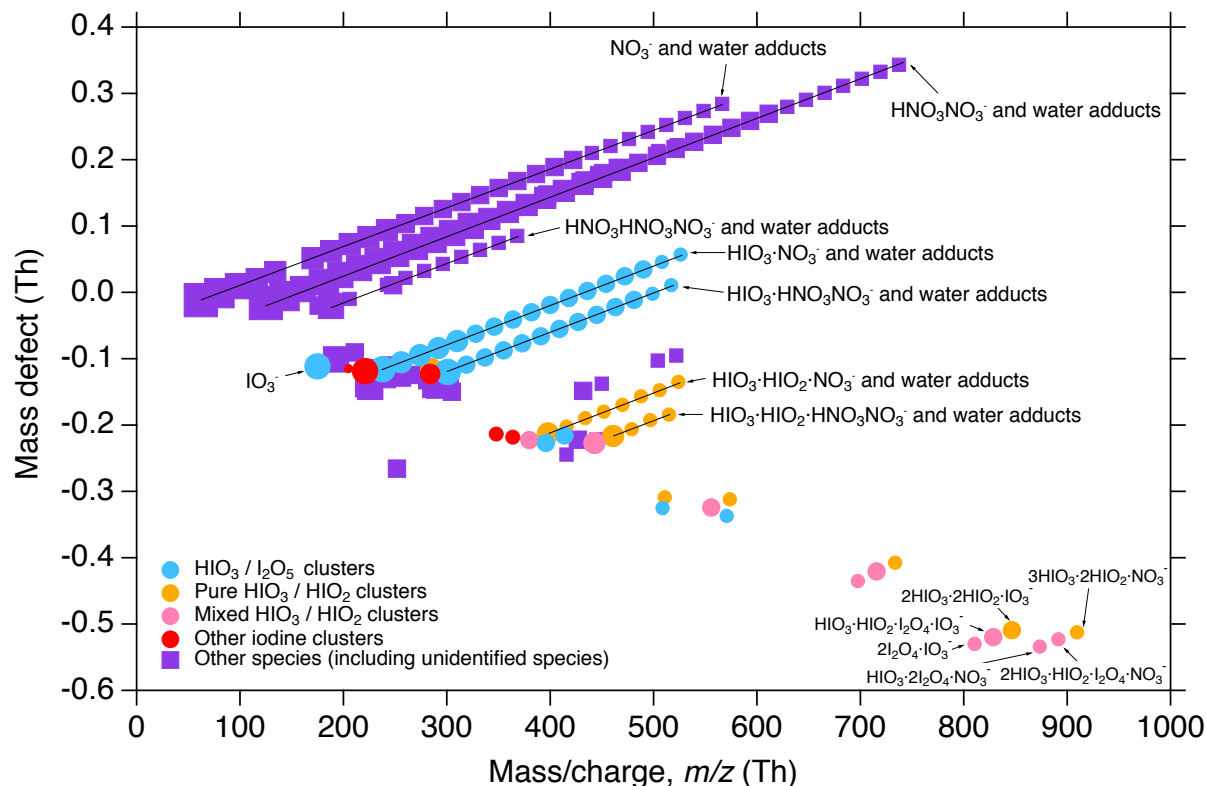


**Fig. S2. Iodine oxoacid and oxide production versus I<sub>2</sub>.** (A) measured concentrations of iodine oxoacids and B) normalized signals of iodine oxides when iodine vapor is adjusted between equilibrium mixing ratios of 0.4 and 4 pptv at fixed experimental conditions of 38 to 42 ppbv O<sub>3</sub>, 34 to 44% RH and +10°C. The I atom production rate is  $4.4 \times 10^4 \text{ cm}^{-3} \text{ s}^{-1}$  to  $3.8 \times 10^5 \text{ cm}^{-3} \text{ s}^{-1}$ . HIO<sub>x</sub> is measured with a nitrate-CIMS, and IO, OIO and I<sub>2</sub> are measured with a bromide-CIMS. The lines are power-law fits to the HIO<sub>3</sub> and HIO<sub>2</sub> concentrations of the form  $\text{HIO}_x = k \times \text{I}_2^n$ , with fitted values for  $n$  of  $(1.07 \pm 0.04)$  and  $(0.51 \pm 0.04)$ , respectively, and the power-law fits to the IO and OIO normalized signals of the form  $\text{IO}_x = k \times \text{I}_2^n$ , with fitted values for  $n$  of  $(0.59 \pm 0.03)$  and  $(0.81 \pm 0.02)$ , respectively.

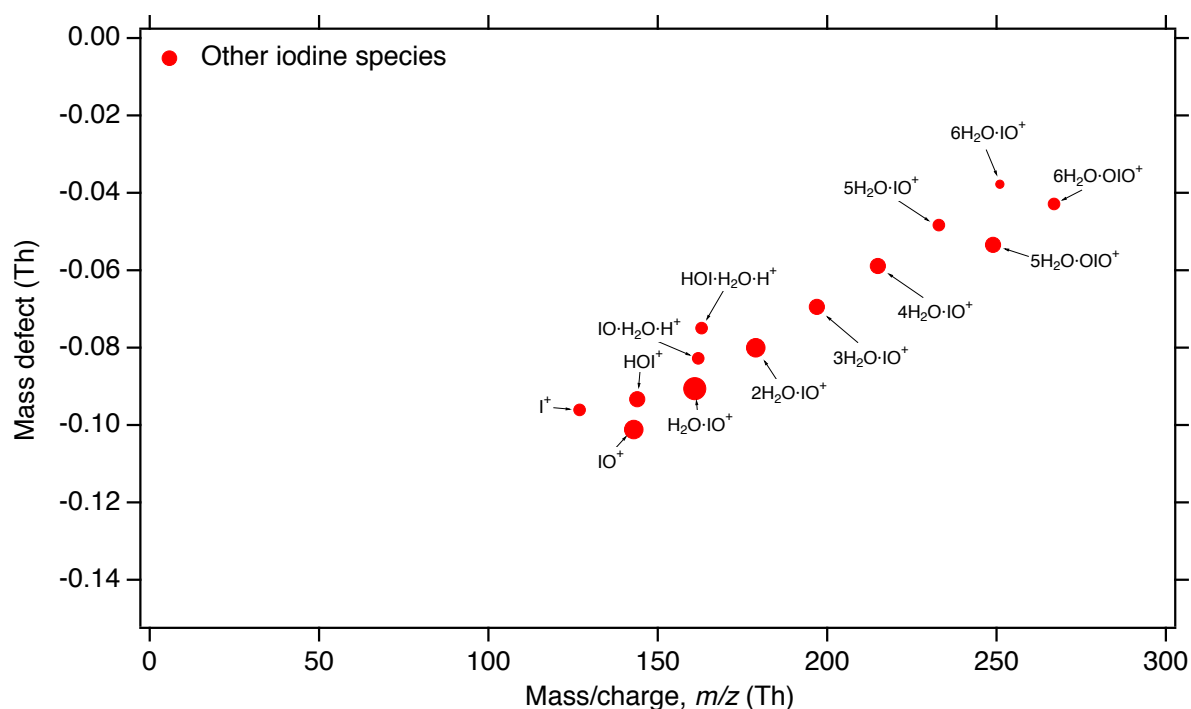




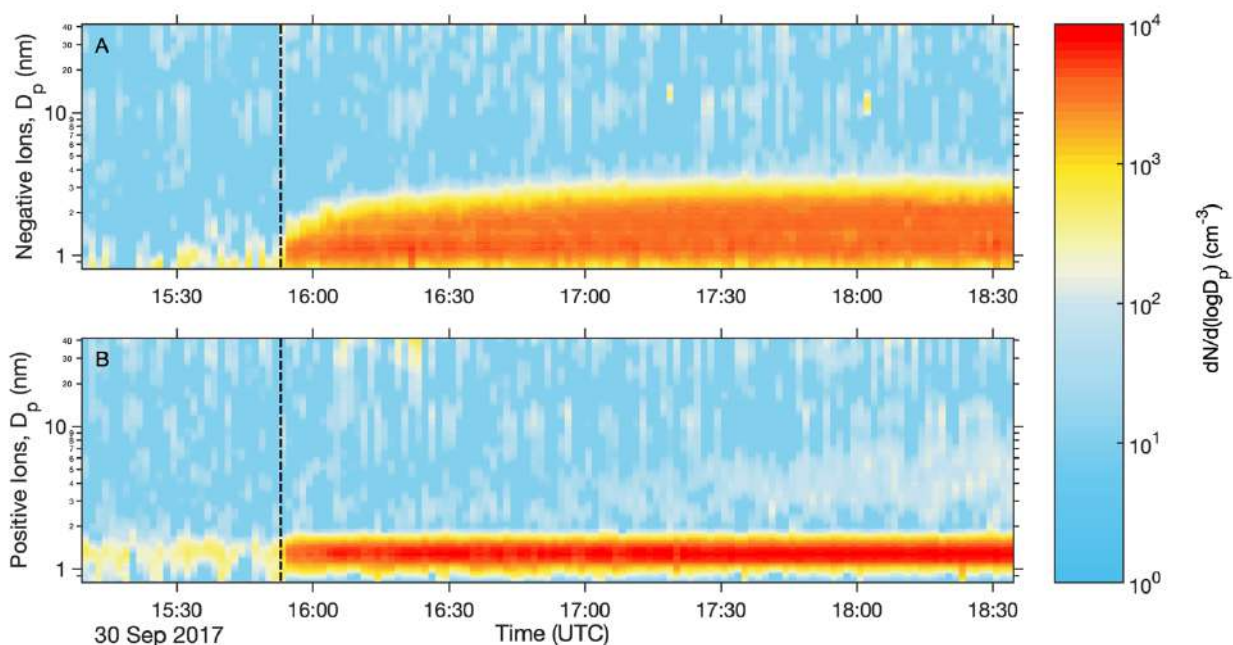
**Fig. S3. Ion-induced iodic acid nucleation.** Cluster mass defect versus  $m/z$  of negatively charged clusters containing up to twelve iodine atoms, measured with the APiTOF(-). These data are the same as in Fig. 2A but extended to a higher mass range. The experimental conditions are 36 ppbv  $O_3$ , 40% RH,  $+10^\circ C$ , 168 pptv  $I_2$  and  $1.5 \times 10^7$  I atoms  $cm^{-3} s^{-1}$ . Blue circles indicate clusters containing only  $HIO_3$  and  $I_2O_5$ . Orange circles indicate clusters containing only  $HIO_3$  and  $HIO_2$ . Pink circles indicate clusters containing  $HIO_3$ ,  $HIO_2$ ,  $I_2O_5$  and  $I_2O_4$ . Red circles indicate other iodine-containing neutral clusters. Purple squares indicate species without iodine, of which some are unidentified. The four sizes of symbol, from smallest to largest, indicate the signal strength (counts per second, cps) on a logarithmic scale: 1 ( $< 0.1$  cps), 2 (0.1-1 cps), 3 (1-10 cps) and 4 ( $> 10$  cps).



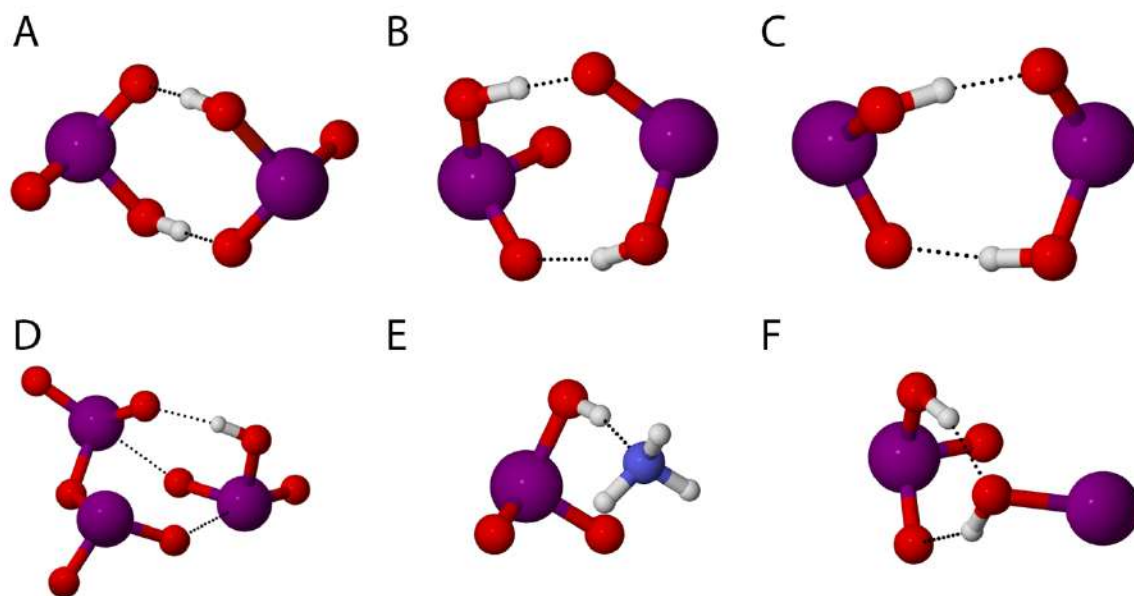
**Fig. S4. Neutral nucleation of iodine oxoacids.** Cluster mass defect versus  $m/z$  of neutral clusters containing up to five iodine atoms, measured with the nitrate-CIMS (preceded by an ion filter). These data are the same as in Fig. 2B but showing all identified peaks without summing over water molecules or charger ions. The experimental conditions are 46 ppbv  $O_3$ , 43% RH,  $+10^\circ C$ , 49 pptv  $I_2$  and  $2.4 \times 10^5$  I atoms  $cm^{-3}s^{-1}$ . Blue circles indicate clusters containing only  $HIO_3$  and  $I_2O_5$ . Orange circles indicate clusters containing only  $HIO_3$  and  $HIO_2$ . Pink circles indicate clusters containing  $HIO_3$ ,  $HIO_2$ ,  $I_2O_5$  and  $I_2O_4$ . Red circles indicate other iodine-containing clusters. Purple squares indicate species without iodine, of which some are unidentified. The diagonal bands indicate clusters with sequential addition of up to 34 water molecules. The six sizes of symbol, from smallest to largest, indicate the signal strength (counts per second, cps) on a logarithmic scale: 1 (<1 cps), 2 (1-10 cps), 3 (10-100 cps), 4 (100-1000 cps), 4 (1000-10000 cps), and 6 (10000 cps).



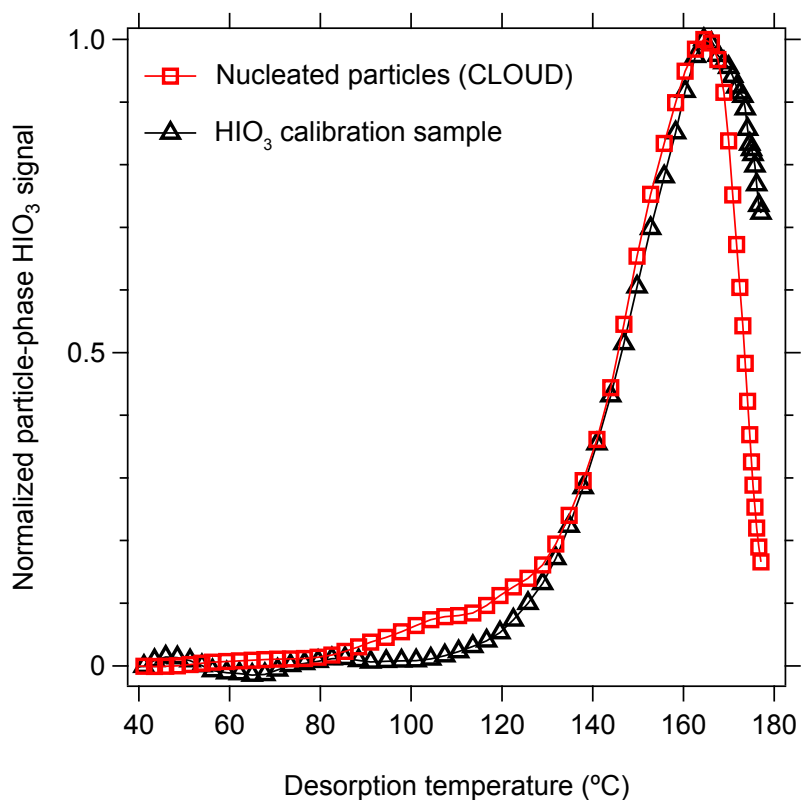
**Fig. S5. Positively charged iodine clusters during a nucleation event.** Cluster mass defect versus  $m/z$  of positively charged clusters during a nucleation event, measured with the APiTOF(+). The experimental conditions are 40 ppbv  $O_3$ , 34% RH and  $+10^\circ C$ , 1.2 pptv  $I_2$  and  $1.1 \times 10^5$  I atoms  $cm^{-3}s^{-1}$ . Red circles represent clusters containing one iodine atom. No positively charged clusters were observed containing more than one iodine atom. The four sizes of symbol, from smallest to largest, indicate the signal strength (counts per second, cps) on a logarithmic scale: 1 (< 0.1 cps), 2 (0.1-1 cps), 3 (1-10 cps) and 4 (>10 cps).



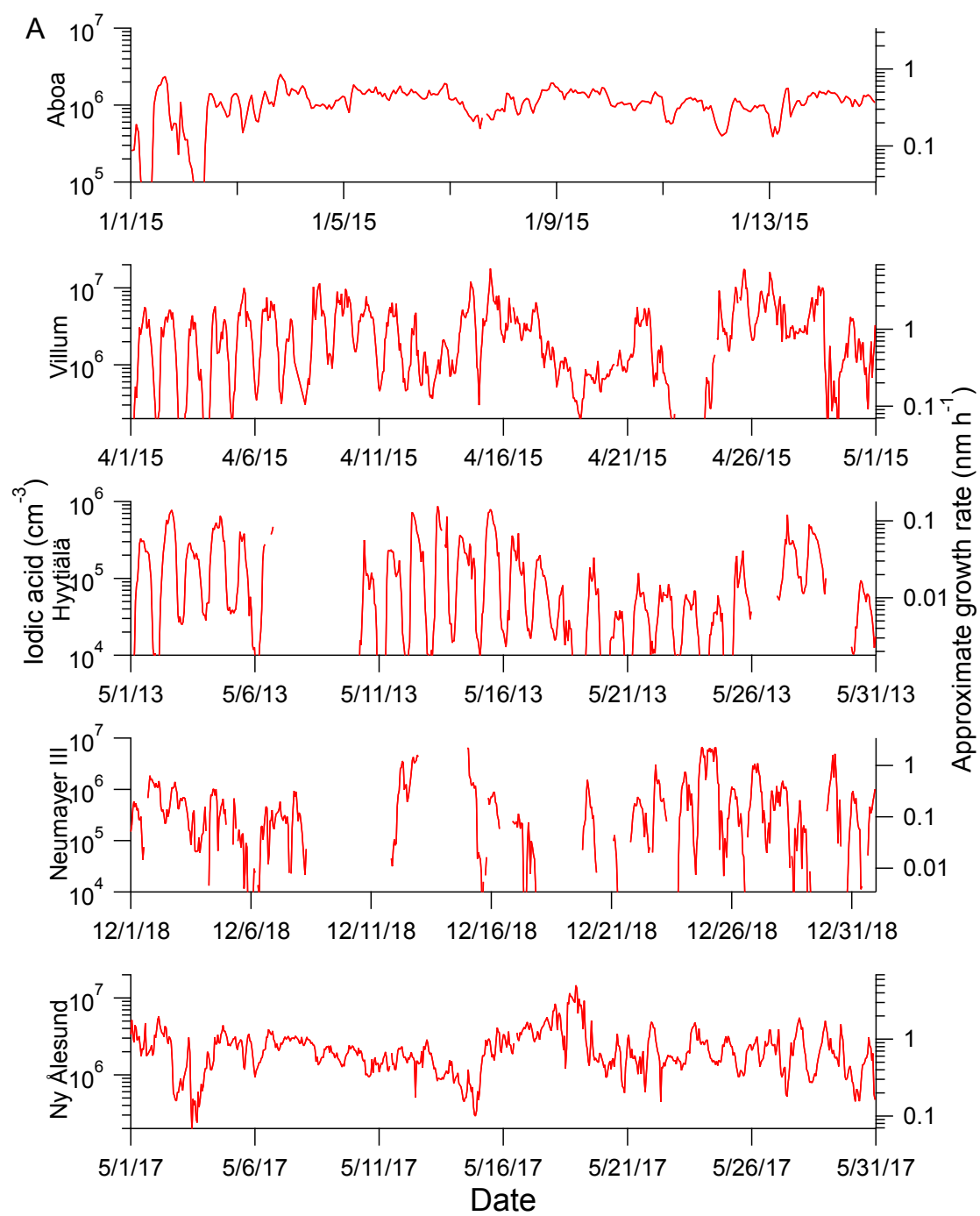
**Fig. S6. Evolution of charged particles during a nucleation event.** Evolution of the size distribution of (A) negative and (B) positive charged particles measured with the NAIS (Neutral cluster and Air Ion Spectrometer) during a nucleation event. The experimental conditions are 40 ppbv  $\text{O}_3$ , 34% RH,  $+10^\circ\text{C}$ , 1.2 pptv  $\text{I}_2$  and  $1.1 \times 10^5$  I atoms  $\text{cm}^{-3} \text{s}^{-1}$ , and zero beam. The vertical dashed line represents the start of the experiment, initiated by turning off the electric field in the chamber and switching on the green light. Small ions of both polarities then build up due to galactic cosmic ray ionization, but nucleation and growth only take place for negative ions. Under these experimental conditions, the negative particles are almost completely neutralized by charge recombination by the time they reach 3 nm, and they continue to grow as mainly neutral particles.



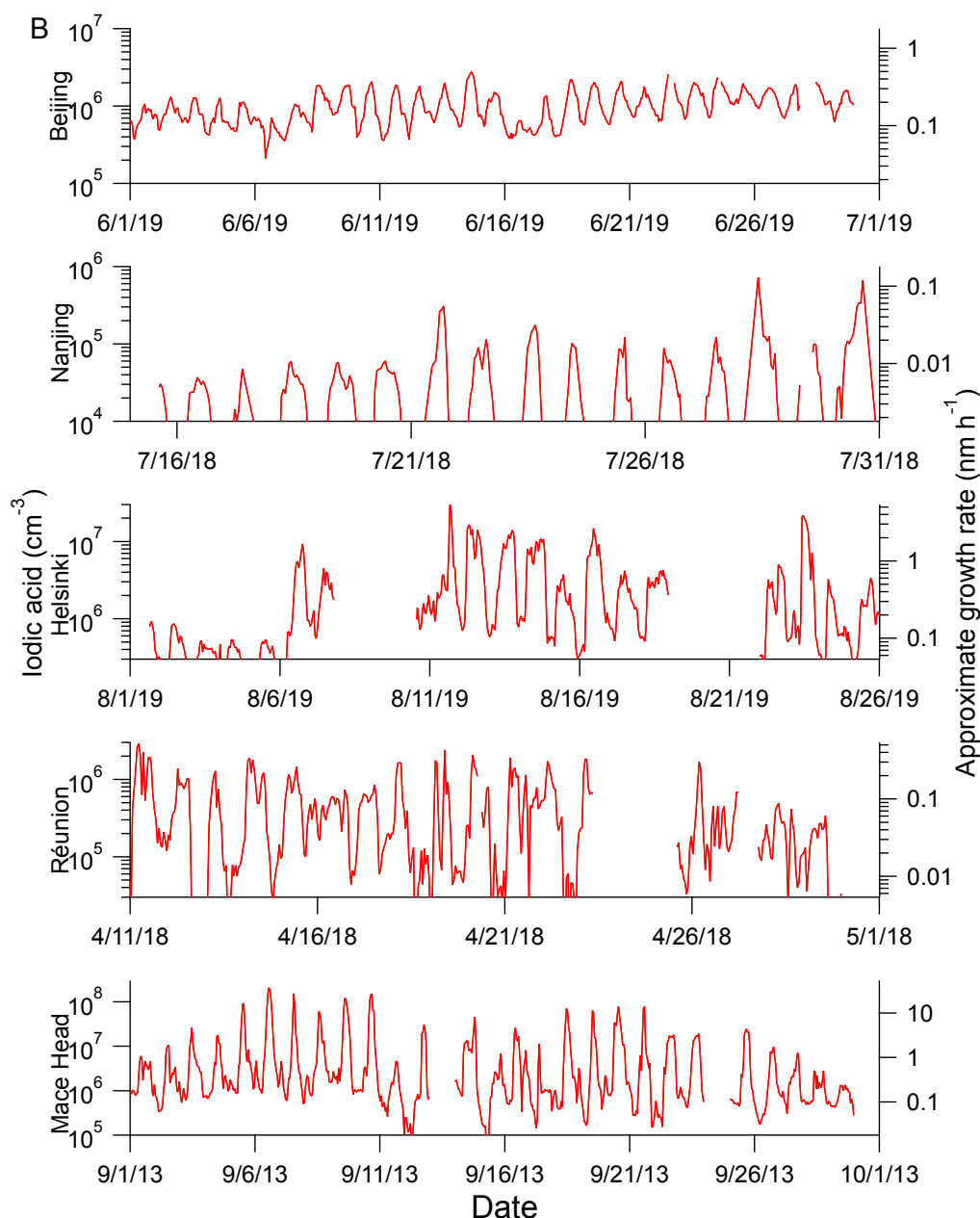
**Fig. S7. Lowest free energy neutral dimers containing iodine oxoacids.** The lowest free energy cluster geometries for (A)  $\text{HIO}_3 \cdot \text{HIO}_3$ , (B)  $\text{HIO}_3 \cdot \text{HIO}_2$ , (C)  $\text{HIO}_2 \cdot \text{HIO}_2$ , (D)  $\text{I}_2\text{O}_5 \cdot \text{HIO}_3$ , (E)  $\text{HIO}_3 \cdot \text{NH}_3$ , and (F)  $\text{HIO}_3 \cdot \text{HOI}$ . The colored balls indicate atoms of iodine (purple), oxygen (red), hydrogen (white), or nitrogen (blue). The clusters are optimized at the  $\omega\text{B97X-D/aug-cc-pVTZ-PP}$  level at 283.15 K. The cluster free energies are summarized in table S3. The most stable dimer with  $\text{HIO}_3$  is  $\text{HIO}_3 \cdot \text{HIO}_2$  (dimer B).



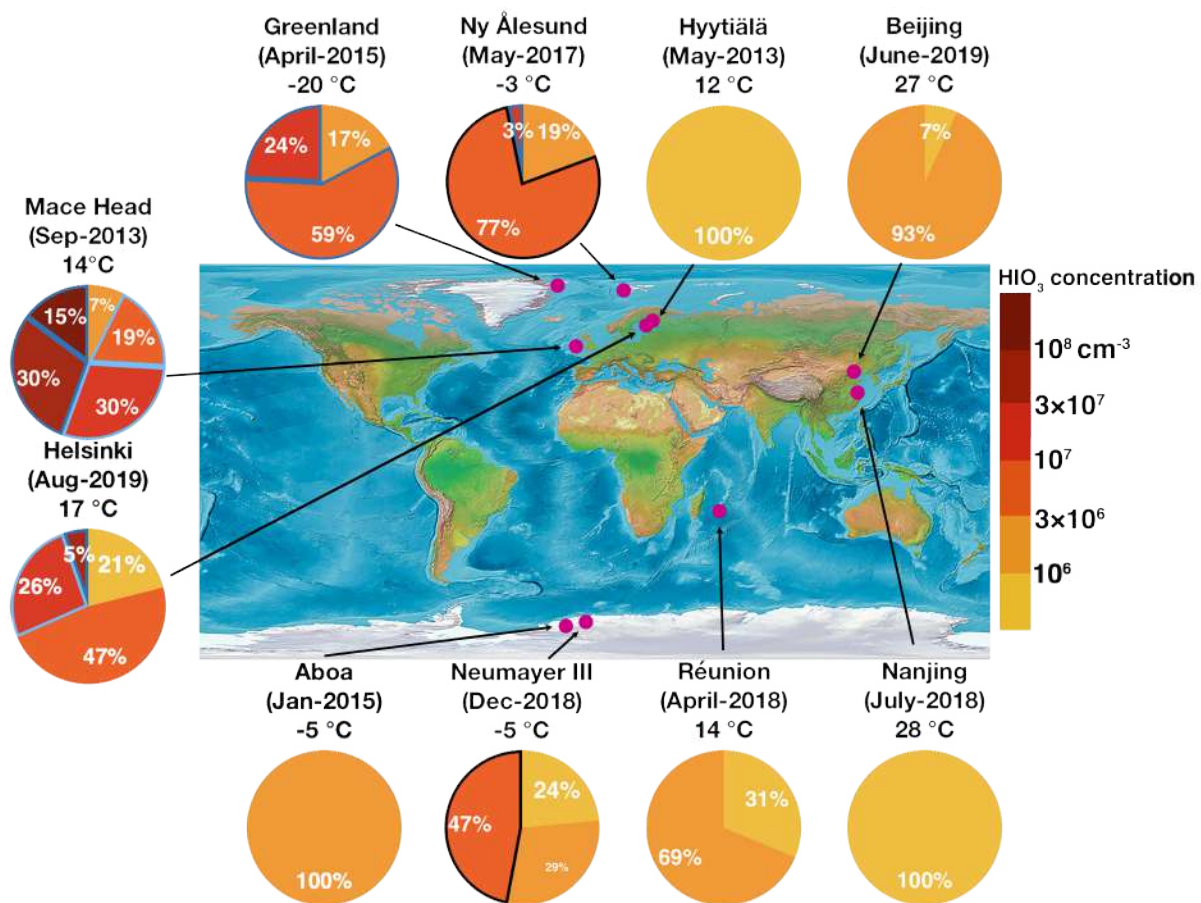
**Fig. S8. Comparison of FIGAERO thermograms for iodic acid.** Normalized Br<sup>-</sup>-FIGAERO thermograms of evaporated HIO<sub>3</sub> samples obtained from an iodine oxoacid particle formation experiment in CLOUD (red squares) and from a nebulized pure HIO<sub>3</sub> calibration sample in the laboratory (black triangles). The two thermograms agree well.







**Fig. S9. Iodic acid measurements at diverse sites.**  $\text{HIO}_3$  concentrations measured at various locations, comprising cities (Beijing and Nanjing, China; Helsinki, Finland), boreal forest (Hyytiälä, Finland), polar regions (Ny Ålesund, Svalbard; Villum research station, Greenland; Aboa and Neumayer III, Antarctica), a coastal marine site (Mace Head, Ireland) and a high altitude site on a tropical island (Réunion, Indian Ocean, at 2160 m). The left axis shows the  $\text{HIO}_3$  concentrations, and the right axis shows the approximate expected particle growth rates based on the measurements reported in Fig. 1B. Our growth rates at  $-10^\circ\text{C}$  are used to estimate those of all the sites with monthly average temperature below  $0^\circ\text{C}$  (Ny Ålesund, Villum research station, Aboa and Neumayer III). Growth rates that we measure at  $+10^\circ\text{C}$  are used for all other sites (Nanjing, Beijing, Helsinki, Hyytiälä, Mace Head and Réunion).



**Fig. S10. Frequency of daily maxima of iodic acid at diverse sites.** Pie charts showing the percentage of days where the daily  $\text{HIO}_3$  maxima fall into the indicated range (evaluated for one-hour-averaged data). Each pie chart represents 2-4 weeks' data at the location, date and mean temperature indicated. Additional information on the sites is provided in section 3. Sectors outlined by light blue and dark blue lines indicate that iodine oxoacid particle formation is expected to be dominated by ion-induced or neutral nucleation, respectively. Sectors outlined by a black line indicate comparable ion-induced and neutral nucleation rates. Sectors without any outline indicate that the expected nucleation rates are below  $0.01 \text{ cm}^{-3} \text{ s}^{-1}$ . The systematic uncertainty between  $\text{HIO}_3$  measurements at different sites is estimated to be a factor three. All measurements are above the  $\text{HIO}_3$  detection limit of the instruments.